

Hydrous wadsleyite crystal structure up to 32 GPa

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ABSTRACT

Hydroxylation of wadsleyite, β -(Mg,Fe)₂SiO₄, is associated with divalent cation defects and well known to affect its physical properties. However, an atomic-scale understanding of the defect structure and hydrogen bonding at high pressures is needed to interpret the influence of water on the behavior of wadsleyite in the mantle transition zone. We have determined the pressure evolution of the wadsleyite crystal symmetry and structure, including all O···O interatomic distances, up to 32 GPa using single-crystal X-ray diffraction on two well-characterized, Fe-bearing (Fo₉₀) samples containing 0.25(4) and 2.0(2) wt% H₂O. Both compositions undergo a pressure-dependent monoclinic distortion from orthorhombic symmetry above 9 GPa, with the less hydrous sample showing a larger increase in distortion at increased pressures due to the difference in compressibility of the split M3 site in the monoclinic setting arising from preferred vacancy ordering at the M3B site. Although hydrogen positions cannot be modeled from the X-ray diffraction data, the pressure evolution of the longer O1···O4 distance in the structure characterizes the primary hydrogen bond length. We observe the hydrogen-bonded O1···O4 distance shorten gradually from 3.080(1) Å at ambient pressure to about 2.90(1) Å at 25 GPa, being still much longer than is defined as strong hydrogen bonding (2.5–2.7 Å). Above 25 GPa and up to the maximum pressure of the experiment at 32.5 GPa, the hydrogen-bonded O1···O4 distance decreases no further, despite the fact that previous spectroscopic studies have shown that the primary O-H stretching frequencies continuously drop into the regime of strong hydrogen bonding (<3200 cm⁻¹) above ~15 GPa. We propose that the primary O1-H···O4 hydrogen bond in wadsleyite becomes highly nonlinear at high pressures based on its deviation from frequency-distance correlations for linear hydrogen bonds. One possible explanation is that the hydrogen position shifts from being nearly on the long O1-O4 edge of the M3 site to a position more above O1 along the *c*-axis.

Keywords: Wadsleyite, mantle transition zone, hydrogen bond, water

INTRODUCTION

The nominally anhydrous minerals (NAMs) wadsleyite and ringwoodite can incorporate 1–2 wt% H₂O into their crystal structures as hydroxyl groups at the elevated pressures and temperatures of the Earth's mantle transition zone (e.g., Smyth 1987; Fei and Katsura 2020), which suggests that the Earth's mantle could play a dynamic role in the evolution of surface water reservoirs over geologic time (Dong et al. 2021). Although no samples of terrestrial wadsleyite have been reported, the recent discovery of hydrous ringwoodite containing ~1.5 wt% water, found as a natural inclusion in a diamond (Pearson et al. 2014), provides evidence that the transition zone is locally hydrated. Mapping regional-scale hydration from seismological data (e.g., Karato 2011; Wang et al. 2018) requires knowledge of the effects

of hydration on the physical properties of NAMs at transition zone pressures and temperatures. For example, theory and experiments show that the incorporation of even small quantities of H₂O significantly influences the elastic properties of wadsleyite governing the velocity of seismic waves (e.g., Tsuchiya and Tsuchiya 2009; Mao et al. 2011; Buchen et al. 2018; Gwanmesia et al. 2020). As the elastic properties of Earth materials are ultimately controlled by interatomic potentials, knowledge of the crystal structure and evolution of hydrogen bonding in wadsleyite is useful in predicting the physical properties expected for hydrous and anhydrous regions of the upper transition zone.

The crystal structure of wadsleyite is unique among the Mg₂SiO₄ polymorphs in possessing a sorosilicate group (Si₂O₇). The structure consists of three octahedral metal cation sites (M) occupied by Mg²⁺, Fe²⁺, or Fe³⁺, one tetrahedral site (T) occupied by Si⁴⁺ or Fe³⁺, and four distinct oxygen sites (Horiuchi and Sawamoto 1981; Smyth et al. 1997) (Online Materials¹ Fig. S1). The first step in understanding the impacts of hydration on the physical and mechanical properties of wadsleyite and other NAMs is

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determining where hydrogen is structurally accommodated and charge-balanced by associated cation defect vacancies. Smyth (1987) recognized that the O1 site of wadsleyite is a favorable location for H based on the electrostatic potential as this site is coordinated to five Mg sites but no Si and is therefore underbonded. Shortly thereafter, McMillan et al. (1991) observed OH absorption bands in wadsleyite using Fourier transform infrared (FTIR) spectroscopy on wadsleyite crystals that had been fortuitously hydrated during synthesis by dehydration of the pyrophyllite pressure medium.

The pleochroic behavior of OH absorbance observed in polarized FTIR spectra (Jacobsen et al. 2005) showed that the primary OH bands at 3300–3400 cm^{-1} can be explained by H at the O1 site, with protonation of the unshared O1–O4 edge of the M3 site, where most cation vacancies were observed. Jacobsen et al. (2005) also noted that some protonation may occur along the shorter O1–O4 edge of the M3 site that is shared with the M2 site, as well as possibly along the O1–O3 edge of the M3 site. Similar sites were proposed from room-pressure XRD and high-pressure FTIR spectroscopy measurements by Deon et al. (2010). Sano-Furukawa et al. (2011) refined the H positions in wadsleyite using powder neutron diffraction data and confirmed that the dominant hydrogen site is located close to O1 and along the longer of the two distinct O1–O4 edges of the M3 site, with minor occupation along the O1–O3 edge of the M3 site. Using single-crystal neutron diffraction data, Purevjav et al. (2016) confirmed that the major H site is along the longer and unshared O1–O4 edges of the M3 site with an O–H distance of ~ 1 Å and slightly nonlinear hydrogen bond angle of 171° . Whereas the longer O1–O4 edges of M3 site appear to host the majority of H in wadsleyite, realistically, there are also many minor sites, as evidenced by up to 15 observed O–H stretching frequencies, which can tentatively be assigned to other sites in the structure based on frequency–distance correlations (Libowitzky 1999; Kohn et al. 2002).

In silicates and oxides, hydrogen bonds form when a H atom, bonded to a donor O atom via a polar covalent bond, interacts electrostatically with a nearby acceptor oxygen (Emsley 1980). Hydrogen bonds have both a characteristic bond length, $d(\text{O}\cdots\text{O})$, defined by the distance between the donor and acceptor O atoms along O–H \cdots O, and a characteristic O–H stretching frequency, which is correlated to hydrogen-bond strength, assuming the hydrogen bond angle along O–H \cdots O is close to 180° (Libowitzky 1999). Typically, hydrogen bonds form an asymmetric potential well between donor and acceptor O atoms, where the O–H bond is shorter than the H \cdots O distance. Although hydrogen bonds in minerals have a poorly defined upper limit beyond ~ 3 Å, strong hydrogen bonds are defined as having $d(\text{O}\cdots\text{O})$ of 2.5–2.7 Å (Libowitzky 1999). On compression, if $d(\text{O}\cdots\text{O})$ shortens, the O–H bond lengthens due to the increased attractive force of the acceptor oxygen, weakening the O–H bond and lowering the O–H stretching frequency (Libowitzky 1999). Whereas the O–H stretching frequencies in most hydrous minerals are observed at 3200–4000 cm^{-1} , strong hydrogen bonds have stretching frequencies in the range of 1600–3200 cm^{-1} .

High-pressure FTIR spectroscopy studies of wadsleyite have shown that the primary O–H stretching bands associated with hydrogen bonding along the O1–O4 edge at 3300–3400 cm^{-1}

shift strongly with increasing pressure to lower frequencies of about 3150–3250 cm^{-1} at 16–18 GPa, the maximum pressure of those experiments (Deon et al. 2010; Yang et al. 2014). This is consistent with earlier findings of Kleppe et al. (2001), who reported Raman spectroscopy measurements of hydrous Fe-bearing wadsleyite up to approximately 50 GPa and found that the 3300–3400 cm^{-1} band became a very broad and weak feature at frequencies below 3000 cm^{-1} above approximately 40 GPa. Because O–H stretching frequencies in hydrous wadsleyite drop below ~ 3200 cm^{-1} above 15 GPa, they enter the regime of strong hydrogen bonding as defined for linear hydrogen bonds (Libowitzky 1999). Since the primary hydrogen bond in wadsleyite, O1–H \cdots O4, is long at room pressure with $d(\text{O1}\cdots\text{O4}) = 3.1$ Å (Jacobsen et al. 2005), a decrease of O–H stretching frequencies to below 3200 cm^{-1} above 15 GPa would imply a drastic compression of the O1–O4 edge if the hydrogen bond remains nearly linear. One objective of this study is to determine all the interatomic O \cdots O distances in the wadsleyite structure up to 32 GPa in part to test whether or not spectroscopic indications of O–H stretching frequencies below ~ 3200 cm^{-1} are consistent with conventionally defined strong hydrogen bonds with $d(\text{O}\cdots\text{O}) < 2.7$ Å.

Another consequence of the hydration of wadsleyite is a change in symmetry from orthorhombic (*Imma*) to monoclinic (*I2/m*), which is defined by the deviation of the unit-cell β -angle from 90.0° . This was first recognized by Smyth et al. (1997) and subsequently observed, to varying degrees, by other studies (e.g., Jacobsen et al. 2005; Holl et al. 2008). The breaking of orthorhombic symmetry in hydrous wadsleyite is believed to result from the unequal distribution of vacancies at M3 sites, resulting in two non-equivalent M3A and M3B sites (Smyth et al. 1997). A Raman spectroscopy study of Fe-bearing wadsleyite with 2.4 wt% water by Kleppe et al. (2006) inferred monoclinic symmetry at high pressure and suggested, based on Raman spectra, that the monoclinic angle might increase with pressure. In contrast, Holl et al. (2008) evaluated the β -angle in wadsleyite containing 1.66 wt% H_2O up to 10 GPa but did not observe a pressure-induced change in β -angle. Determining the structure and symmetry of hydrous wadsleyite at pressures exceeding 10 GPa is necessary to more adequately understand the physical properties of hydrous wadsleyite in the transition zone.

To date, only one experimental study (Hazen et al. 2000) has examined the structure of wadsleyite at high pressures ($P_{\text{max}} = 10.1$ GPa) from single-crystal X-ray diffraction data, but the water content of the samples in that study was not determined. Therefore, we have undertaken high-precision structure refinements of well-characterized hydrous wadsleyite single crystals under quasi-hydrostatic compression up to 35 GPa. To directly compare the effects of water, one sample contains ~ 2.0 wt% H_2O while the other sample contains ~ 0.25 wt% H_2O . Both samples approximate the Fe content of mantle olivine (Fo_{90}). Our two main objectives are: (1) to evaluate the crystal symmetry evolution with pressure and (2) to determine the pressure-dependence of all interatomic O \cdots O distances in the structure to compare hydrogen-bonded vs. non-hydrogen bonded O–O edges for comparison to previously observed O–H stretching frequencies at high pressures. The results of this study help to explain the modified physical properties of hydrous wadsleyite in the transition zone.

MATERIALS AND METHODS

Sample synthesis and characterization

This study examines two different hydrous Fe-bearing wadsleyite samples from synthesis run Z0570 containing 2.0(2) wt% H₂O and from run Z0902 containing 0.25(4) wt% H₂O (Chang et al. 2015). Both samples were synthesized using the 5000-ton, multi-anvil press at Bayerisches Geoinstitut, University of Bayreuth, Germany. For run Z0570, starting materials of San Carlos olivine plus liquid water were welded into a Pt capsule, compressed to 18 GPa, and heated to a peak temperature of 1400 °C for several minutes before annealing at 1100 °C for 30 min. Run products from Z0570 consisted of both wadsleyite and ringwoodite. For run Z0902, San Carlos olivine was used without the addition of water and compressed to 11 GPa for annealing at 1100 °C (Chang et al. 2015). No phase other than wadsleyite was observed in the run product of Z0902. Electron microprobe analyses give values of Mg/(Mg+ΣFe) of 0.89 and 0.90 for run Z0570 and run Z0902, respectively (Chang et al. 2015). FTIR spectra for both samples are presented in Chang et al. (2015) and indicate 1.4 wt% H₂O and 0.25 wt% H₂O in run Z0570 and Z0902, respectively, using the calibration of Deon et al. (2010). The water content of Z0570 was also measured at two secondary ion mass spectrometry laboratories by Chang et al. (2015), resulting in 1.93(22) wt% H₂O and 2.07(24) wt% H₂O. Since FTIR methods are less reliable for very high water contents in wadsleyite, we used 2.0(2) wt% H₂O for the more hydrous sample (Z0570) and 0.25(4) wt% H₂O for the less hydrous sample (Z0902).

Synchrotron Mössbauer spectroscopy

To determine the oxidation states of iron, time-domain synchrotron Mössbauer spectroscopy (SMS) was conducted on beamline 3-ID-B of the Advanced Photon Source (APS), Argonne National Laboratory. One ~150 μm diameter grain was selected from each synthesis run. A Si (111) double crystal monochromator and a 4-bounce inline high-resolution monochromator (Toellner 2000) was used to filter the X-ray energy to a bandwidth of 1 meV at 14.4125 keV. The monochromatic X-rays were focused into a 15 μm diameter beam using a Kirkpatrick-Baez type mirror. The nuclear delay signal was recorded in the 21–128 ns time window between the X-ray pulses generated by the synchrotron. Time-domain Mössbauer spectra were fit using CONUSS 2.2.0 (Sturhahn 2000) to get the hyperfine parameters of Fe and the ferric-to-ferrous ratio for each sample. A two-doublet model was used to fit each spectrum, with one doublet each for Fe²⁺ and Fe³⁺, as it gave the best statistical fit. The isomer shift of Fe²⁺ was fixed at 1.04 mm/s relative to stainless steel based on previous Mössbauer data (Chang et al. 2015).

Sample loading environment and pressure determination

High pressures were achieved by loading samples into one of two short symmetric type diamond-anvil cells (DACs) with either a 78° or 81° opening angle. Sample chambers of 180 μm diameter were laser ablated into rhenium gaskets that were pre-indentured to ~40 μm thickness. For each sample, a single crystal of wadsleyite ~30 μm in diameter was parallel-polished down to a thickness of ~10 ± 2 μm. Each sample was loaded into a DAC equipped with Boehler-Almax diamond anvils with 300 μm culets. Two ruby spheres were placed inside each cell as pressure standards, with one ruby placed close to the center of the sample chamber to serve as the primary pressure standard and the other ruby placed near the edge of the sample chamber to monitor the possible development of pressure gradients. Ruby pressures were calculated from the shift of the R₁ ruby fluorescence line using the quasi-hydrostatic calibration from Jacobsen et al. (2008). Neon was loaded using the COMPRES-GSECARS gas loading system (Rivers et al. 2008) to serve as a quasi-hydrostatic pressure-transmitting medium.

X-ray diffraction experiments

Single-crystal X-ray diffraction (XRD) data were collected on beamline 13-BM-C of the APS. The incident X-ray beam was monochromatized with a silicon (311) crystal to 0.434 Å incident wavelength with 1 eV energy bandwidth and focused to a 15 × 15 μm² spot. A Pilatus 1M (Dectris) detector with 1 mm thick silicon sensor was used to record the diffraction patterns. The sample-to-detector distance and the tilt of the detector were calibrated using LaB₆ (Zhang et al. 2017). At each pressure, step scans were collected in steps of 1° at 1 s⁻¹ in the φ direction using a detector with two collecting positions. The collecting positions were moved to be positioned 20° apart in 2θ angle. Diffraction intensity data for crystal structure refinements were collected at ambient conditions before gas loading and at each subsequent pressure step. High-pressure data were collected by compressing the DAC with a gas membrane cell at pressure intervals

of 1–1.5 GPa. After increasing the pressure at each pressure step, the sample was allowed to sit for 10–15 min before data collection to stabilize the pressure. The reported pressures are the average of pressures measured before and after XRD data collection. The reported uncertainty is half of the absolute pressure difference before and after XRD collection. When this pressure difference was smaller than 0.1 GPa, 0.1 GPa was used as the pressure uncertainty to reflect the linewidth of the ruby fluorescence spectrum.

Diffraction intensities and unit-cell parameters were obtained using the APEX3 software (Bruker). SHELXL (Dolomanov et al. 2009; Sheldrick 2015), with the Olex2 general user interface was used to refine the crystal structure using peak intensities reduced by the APEX3 software. All crystal structures were refined using intensity data collected from both of the collecting positions of the detector. Two models were used to refine the crystal structure: an orthorhombic (*Imma*) model from Jacobsen et al. (2005) and a monoclinic (*I2/m*) model from Smyth et al. (1997). The difference between them is that Smyth et al. (1997) split the M3 site into two non-equivalent sites denoted M3A and M3B, and the O4 site is split into O4A and O4B, where O4A is bonded to M3A and O4B bonded to M3B. Refinement parameters at each pressure step, including the 2θ range, number of reflections, *hkl* range, and merging and internal R-factors are listed in Online Materials¹ Tables S1 and S2.

RESULTS

Mössbauer spectroscopy

Ambient-pressure SMS data for the more hydrous wadsleyite (Z0570) show that 24.7(1.1)% of the total Fe was Fe³⁺ with an isomer shift of 0.082(9) mm/s and quadrupole splitting of 0.358(9) mm/s. The quadrupole splitting for Fe²⁺ was 2.638(3) mm/s. For the less-hydrous wadsleyite sample (Z0902), 12.8(1.0)% of the total Fe was Fe³⁺ with an isomer shift of 0.106(17) mm/s and quadrupole splitting of 0.25(2) mm/s. The quadrupole splitting for the Fe²⁺ site was 2.657(4) mm/s (Table 1). The isomer shifts (relative to stainless steel), and quadrupole splitting values reported here for Z0570 are an improvement over those measured by conventional Mössbauer spectroscopy in Chang et al. (2015) because of the improved signal-to-noise ratio in the current data. Chang et al. (2015) did not measure the Fe³⁺ content of Z0902. Combining the current SMS results with the major element analyses and water content determinations from Chang et al. (2015), we obtain the chemical formulas (Mg_{1.638}Fe_{0.145}Fe_{0.047})H_{0.320}SiO₄ for sample Z0570 and (Mg_{1.782}Fe_{0.175}Fe_{0.025})H_{0.041}SiO₄ for sample Z0902. The measured time-domain spectra and calculated energy-domain spectra of the best-fit hyperfine model are shown in Online Materials¹ Figure S2.

Equations of state

Our ambient-pressure diffraction data show that sample Z0902 (0.25 wt% H₂O) has a slight monoclinic distortion, with β = 90.034(18)° and lattice parameters of a₀ = 5.7119(5), b₀ = 11.4785(5), c₀ = 8.282(2) Å, resulting in a calculated unit-cell volume of 543.00(14) Å³. The more hydrous wadsleyite sample, Z0570 (2 wt% H₂O), also has a slight monoclinic distortion, with β = 90.015(3)° and lattice parameters of a₀ = 5.6965(2), b₀ = 11.5431(4), c₀ = 8.269(3) Å, resulting in a calculated unit-cell

TABLE 1. Comparison of hyperfine parameters of the two wadsleyite samples

	Sample Z0570	Sample Z0902
Fe ³⁺ /ΣFe	24.7(1.1)%	12.8(1.0)%
Fe ²⁺ isomer shift (mm/s)	1.04	1.04
Fe ³⁺ isomer shift (mm/s)	0.082(9)	0.106(17)
Fe ²⁺ quadrupole splitting (mm/s)	2.638(3)	2.657(4)
Fe ³⁺ quadrupole splitting (mm/s)	0.358(9)	0.252(20)

volume of $543.7(2) \text{ \AA}^3$. Based on these unit-cell volumes and the detailed compositions given above, the reference density for the more hydrous sample Z0570 is $\rho_0 = 3.5064(13) \text{ g/cm}^3$, and that for the less hydrous sample Z0902 is $\rho_0 = 3.584(1) \text{ g/cm}^3$.

During compression up to about 10 GPa, the β -angle for both samples was essentially invariant. However, above 9 GPa, both samples exhibit a linearly increasing β -angle with pressure, although by markedly different amounts (Fig. 1). Whereas the β -angle in the more hydrous sample (Z0570) only increased by about 0.1° between 10 and 30 GPa, reaching $90.09(2)^\circ$ at the highest pressure of 33 GPa, the β -angle of the less hydrous sample (Z0902) increased by nearly 0.5° , reaching a value of $90.53(4)^\circ$ at the maximum pressure of 35 GPa. The lattice parameters at each pressure refined for both orthorhombic and monoclinic symmetry for both samples are provided in Online Materials¹ Tables S3–S6.

Unit-cell volumes were also calculated using both orthorhombic and monoclinic symmetry for both samples at each pressure (Online Materials¹ Tables S3–S6). We fit the pressure-volume (P - V) data to a third-order Birch-Murnaghan equation of state (BM3-EOS), where K_0 is the reference bulk modulus at ambient pressure, K'_0 is the first pressure derivative of the bulk modulus evaluated at $P = 0$, and V_0 is the reference volume at ambient pressure. The EOS fit was performed with EOSFit7c (Angel et al. 2014). The inverse of the measurement uncertainties was used as weights for least-squares fits. The resulting EOS parameters for the more hydrous sample (Z0570) using orthorhombic symmetry are: $V_0 = 543.19(23) \text{ \AA}^3$, $K_0 = 164.0(2.7) \text{ GPa}$, and $K'_0 = 4.26(23)$. Assuming monoclinic symmetry for Z0902 we obtain $V_0 = 543.09(22) \text{ \AA}^3$, $K_0 = 172.3(2.3) \text{ GPa}$, and $K'_0 = 3.91(15)$. Based on statistical significance, our preferred EOS fit to the more hydrous sample (Z0570) assumes orthorhombic sym-

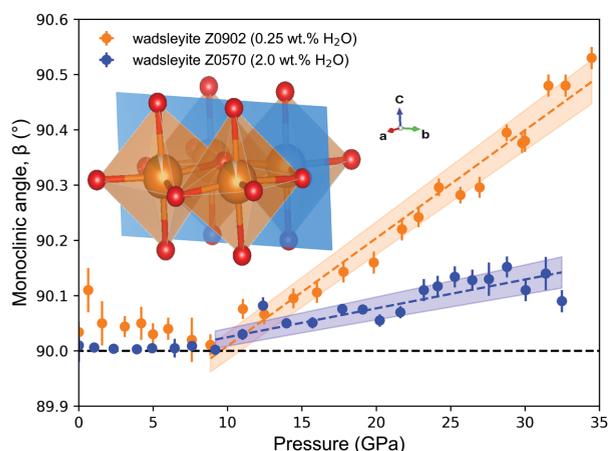


FIGURE 1. Monoclinic β -angle vs. pressure of hydrous wadsleyite. Orange and blue dashed lines are linear fits to the data above 9 GPa for Z0902 and Z0570, respectively. The colored shaded regions indicate the 95% prediction bands. The black dashed line indicates the β -angle of the orthorhombic structure. The inset model by VESTA (Momma and Izumi 2011) shows the local structure around the M3 octahedral site, wherein the blue plane indicates the mirror plane perpendicular to the a -axis for orthorhombic symmetry. The mirror plane is lost when vacancy ordering splits the M3 site into the non-equivalent M3A and M3B sites. (Color online.)

metry and monoclinic symmetry for the less hydrous sample (Z0902), which are plotted in Figure 2. EOS parameters for both compositions in both crystal systems are given in Online Materials¹ Table S7.

Structure refinements

Variation of the lattice parameters with pressure are plotted in Figure 3. A linearized third-order BM-EOS fit was used to determine the linear compressibilities along the crystal axes (Angel et al. 2014). The resulting linear moduli and axial compressibilities are given in Online Materials¹ Table S8. In both samples, the c -axis is the most compressible, and the a - and b -axes have about the same compressibility. However, with increased water content, the c -axis becomes more incompressible, while the a - and b -axes become more compressible. We also evaluated the M-O bond lengths and polyhedral volumes as a function of pressure, which are shown in Online Materials¹ Figures S3 and S4 and listed in Online Materials¹ Table S8. Compression data of all polyhedral volumes are plotted in Figure 4.

The ambient-pressure structure refinements show vacancies only at the M3 sites, similar to what has been observed in previous studies (Smyth 1994; Jacobsen et al. 2005; Holl et al. 2008; Purevjav et al. 2016). For the less hydrous and monoclinic sample (Z0902), the total site occupancies were 100% for M1, M2, and M3A, but 94.9% in M3B, indicative of strongly preferred vacancy ordering onto the M3B site, which has not previously been reported. The refined Fe occupancies at each site in Z0902 were 9.8% in M1, 3.8% in M2, 1.2% in M3A, and 9.4% in M3B. For the very hydrous sample (Z0570), we observed no vacancy ordering between M3A and M3B and used the orthorhombic symmetry, finding the total site occupancies were 100% for M1 and M2, but 87.8% in M3, consistent with there being more hydrogen-related cation vacancies in the more hydrous sample. In Z0570, we refined Fe/Mg as well and found 11.2% Fe in M1, 0.8% in M2, and 6.4% in M3.

The refined site occupancies obtained at room pressure were fixed for the high-pressure structure refinements, except for the M3 site (M3A and M3B), which contains vacancies. However,

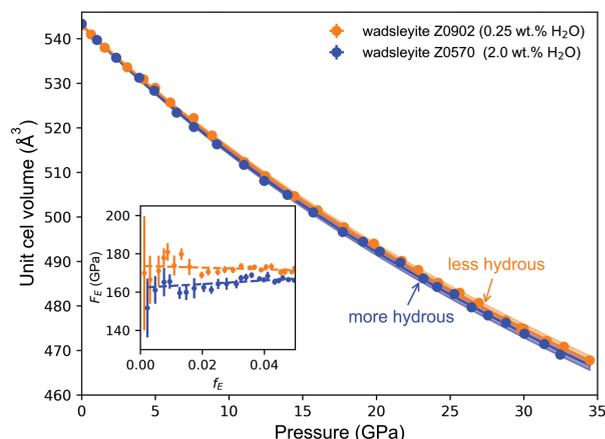


FIGURE 2. Unit-cell volume vs. pressure data for hydrous Fe-bearing wadsleyite. Fitted third-order BM-EOS are shown as solid curves. The blue and orange shaded regions show the 95% prediction band. The inset shows normalized pressure (F_E) vs. Eulerian strain (f_E). (Color online.)

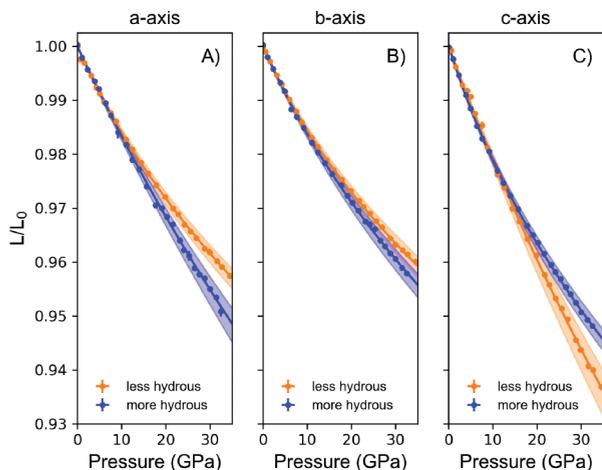


FIGURE 3. Axial compressibility of wadsleyite sample Z0570 with 2.0 wt% H₂O (blue data) and sample Z0902 with 0.25 wt% H₂O (orange data). Fitted third-order BM-EOS fits are shown as solid curves. The shaded regions show the 95% prediction band. (Color online.)

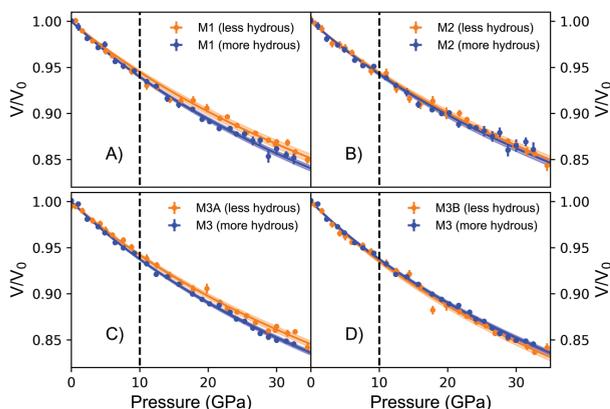


FIGURE 4. Octahedral volume compression of the (a) M1 (b) M2, and (c and d) M3 sites of hydrous wadsleyite sample Z0570 with 2.0 wt% H₂O (blue) and sample Z0902 with 0.25 wt% H₂O (orange). Solid lines show the fitted second-order BM-EOS. The shaded regions show the 95% prediction band. Dashed lines indicate the pressure at which the structure distorts from orthorhombic to monoclinic symmetry. (Color online.)

we did not observe significant changes in vacancy ordering with increasing pressure. Our results also suggest that there is no Fe³⁺ at the T site of slightly hydrous wadsleyite, but up to 4% of the Fe cations may be at the T site of very hydrous wadsleyite, consistent with previous findings (Smyth et al. 2014).

To illustrate how hydrogen potentially affects the compressibility of O-O octahedral edges in the wadsleyite structure, we plot selected O-O distances as a function of pressure in Figure 5. A second-order BM-EOS was used to fit the data with EOSFit7c (Angel et al. 2014). In Figure 5, we also plotted prediction bands instead of confidence intervals to indicate that these are the error bars on quantities calculated from a model, in contrast to error calculated directly from experimental data. The inverse of the measurement uncertainties was used as weights in least-squares fits. The 95% prediction bands of the O1-O4 (short) and O1-O1

edges for more hydrous and less hydrous wadsleyite nearly overlap, indicating very similar compressibility. The long O1-O4 and O1-O3 edges, however, respond differently under compression in the two samples: whereas the long O1-O4 edge is less compressible in the more hydrous wadsleyite sample (Z0570), the O1-O3 edge is less compressible in the less hydrous wadsleyite. Most notably, the longer O1-O4 edge in the more hydrous sample becomes incompressible above ~25 GPa (Fig. 5), but no comparable stiffening was observed in any other O-O distances, except possibly in the O1-O3 edge of the more hydrous sample, which stiffens with a lesser degree of certainty. As stated earlier, the longer O1-O4 octahedral edge is the primary hydrogen bond in wadsleyite (Sano-Furukawa et al. 2011; Purevjav et al. 2016). A comprehensive overview of all O-O distances at the M1, M2, and M3 sites comparing both H₂O compositions is given in Online Materials¹ Figure S5.

DISCUSSION

Ordering in wadsleyite

Our observation that Fe-bearing hydrous wadsleyite distorts from orthorhombic to monoclinic symmetry with an increasing β -angle above 9 GPa could explain earlier observations of new Raman peaks above ~9 GPa in the Raman spectra of hydrous Fe-bearing wadsleyite (Kleppe et al. 2006). Jacobsen et al. (2005) and Holl et al. (2008) found that in pure-Mg wadsleyite there is a correlation between monoclinic distortion and increasing water content. In contrast, in our Fe-bearing samples, our less hydrous sample with 0.25 wt% H₂O actually deviates more from orthorhombic symmetry than the more hydrous sample with ~2 wt% H₂O. To date, the crystal symmetry of hydrous wadsleyite has not been investigated as a function of pressure. Surprisingly, on compression above 9 GPa, the less hydrous sample deviates significantly more from orthorhombic symmetry than the more hydrous sample (Fig. 1). Thus, the symmetry of Fe-bearing hydrous wadsleyite cannot be understood based on water content alone.

In the less hydrous wadsleyite sample, the M3B octahedron has a slightly larger volume than the M3A octahedron. This volume difference is related to M-site vacancies and may be caused by enhanced O-O repulsion in the absence of a central cation (Jacobsen et al. 2005). If the larger polyhedral volume of the M3B site is the result of a higher vacancy concentration as compared to the M3A site, the M3B site can be expected to be more compressible, which is precisely what we observe in the polyhedral moduli (Table 2) and interatomic distances (Online Materials¹ Table S9). Due to the increased compressibility of the M3B site, at pressures exceeding 10 GPa it becomes smaller in volume than the M3A site (Online Materials¹ Tables S12–S13 and Fig. S6), which corresponds to the pressure at which we begin seeing the more distinct symmetry change from orthorhombic to monoclinic. Orthorhombic symmetry in wadsleyite requires the plane perpendicular to the *b*-axis between equivalent M3 sites as a mirror plane (Smyth et al. 1997; Ye et al. 2010, 2011). At low pressure, this plane is almost perpendicular to the *a*-axis but slightly tilted. With increasing pressure, the volume of the M3B octahedron decreases more than that of the M3A octahedron and, at around 10 GPa, drops below the volume of the M3A octahedron (Online Materials¹ Tables S12–S13 and Fig. S6),

FIGURE 5. Compression of selected O···O interatomic distances in hydrous wadsleyite. The structural models are by VESTA (Momma and Izumi 2011). (a) The primary hydrogen bond: in red and orange, long O1-O4 edges on the M3A and M3B sites for sample Z0902 with 0.25 wt% H₂O and monoclinic crystal symmetry. In blue, long O1-O4 edge on the M3 site in sample Z0570 with 2.0 wt% H₂O and orthorhombic crystal symmetry. The dashed vertical line shows the pressure above which the long O1-O4 edge of the M3 site stops shortening above ~25 GPa. (b) Same as panel (a) but showing the short O1-O4 edges of the M3 sites. (c) Other possible hydrogen bonds: O1-O3 of Z0902 and Z0570, with a dashed vertical line to show the pressure above which a possible change in the O1-O3 compression occurs; (d) O1-O1 of Z0902 and Z0570. The distances are normalized to the distance at ambient pressure. Lines show the fitted second-order BM-EOS. Shaded regions show the 95% prediction band. The insets show where the various O-O edges are located in the crystal structure. (Color online.)

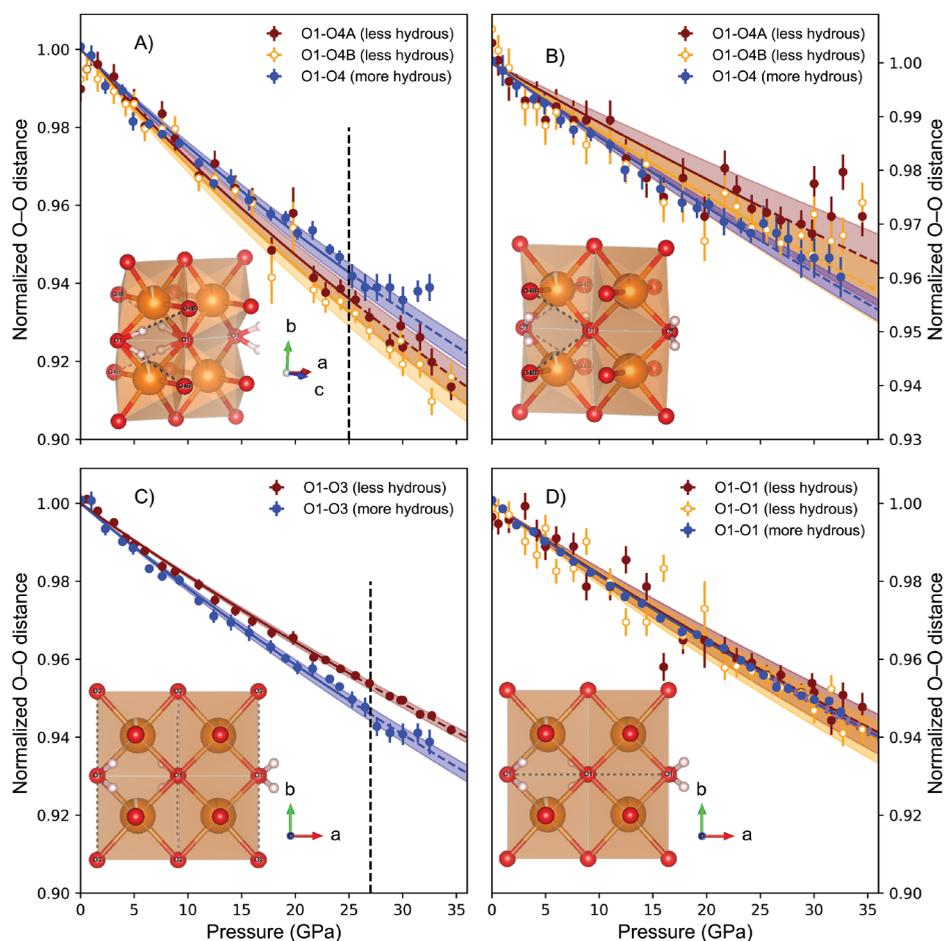


TABLE 2. Compressibility of the coordination polyhedra of wadsleyite with 2.0 wt% H₂O (Z0570) and with 0.25 wt% H₂O (Z0902)

Sample	M1 (GPa)	M2 (GPa)	M3 (GPa)	M3A (GPa)	M3B (GPa)	T (GPa)
Z0570	142(1.9)	150.7(2.4)	134.2(1.5)			314(10)
Z0902	154(3.5)	152.5(3.2)		145.9(3.2)	130.3(2.6)	316(13)

leading the (010) plane to increasing tilt, which breaks the mirror symmetry and causes the β -angle to deviate from 90°. In short, the evolution of the β -angle at high pressures is likely due to the difference in vacancy population of the two non-equivalent M3 sites and their corresponding differences in polyhedral compressibility. We propose that the monoclinic distortion in wadsleyite increases with pressure when there is a significant difference in the vacancy populations or in the Mg-Fe occupancies of the M3A and M3B sites, leading to those sites having different compressibility and thus further distortion from orthorhombic symmetry. The more hydrous wadsleyite sample in our study had no significant difference in the occupancies of the M3 sites, which have similar compressibility, thus the plane perpendicular to the b -axis was retained as a mirror symmetry element.

Our observations of changes in the crystal symmetry at ~9 GPa may shed light on some previous spectroscopic observations. In Fe-free nominally anhydrous wadsleyite, Chopelas (1991) observed

seven Raman mode changes at 9.2 GPa, which the author posited as evidence of a second-order phase transition. Cynn and Hofmeister (1994) reported high-pressure infrared data for Fe-bearing wadsleyite and described a minor structure change at 9 GPa. Kleppe et al. (2006) reported the addition of four new Raman modes in the mid-frequency range at 9 GPa in hydrous Fo₉₀ wadsleyite. If the change in wadsleyite crystal symmetry to monoclinic symmetry above 9 GPa is due to the M3 and O4 sites splitting into two non-equivalent sites, the changes in vibrational modes are likely to result from a distortion of individual octahedra since the affected vibrations are related to MgO₆ octahedra. Our high-pressure structure data could, therefore, potentially reconcile several previously observed changes in vibrational spectra at high pressures.

Influence of hydroxyl groups on the equation of state

The elastic moduli and their pressure derivatives are needed to calculate density and sound velocities at high pressures. For wadsleyite, previous studies have shown that the bulk modulus (K_0) decreases with water content (e.g., Jacobsen et al. 2005) and increases with iron content (e.g., Hazen 1993), as recently reviewed by Buchen et al. (2017) and Wang et al. (2018). In agreement with these trends, our K_0 of ~164 GPa for hydrous Fo₈₉ wadsleyite with 2 wt% H₂O is about 6% lower than dry Fo₉₀ wadsleyite (Li

and Liebermann 2000; Liu et al. 2009). The relationship between water content and the pressure derivative of the bulk modulus (K'_0) shows no obvious trend (e.g., Chang et al. 2015). Holl et al. (2008) suggested that K'_0 should increase with water content because the strong O-O repulsive forces on compression are ultimately a more important factor than the initially larger and more compressible partially vacant M sites. But experiments by Holl et al. (2008) and Chang et al. (2015) found no link between water content and K'_0 . Unfortunately, attempts to resolve this discrepancy have been hampered by a range of derived values for K'_0 that can be dependent on experimental factors, including data density, pressure range, pressure medium, and pressure calibration, among others.

Buchen et al. (2018) obtained $K'_{S0} = 4.13(8)$ for a slightly hydrous wadsleyite (~0.24 wt% water) with Fo_{89} composition using Brillouin spectroscopy, and $K'_{T0} = 4.4(2)$ on the same sample from static compression using the BM3-EOS to $P_{\text{max}} = 20$ GPa with ruby pressures and 10 data points (Buchen et al. 2017), where the subscript S in K'_{S0} refers to the adiabatic bulk modulus and the subscript T in K'_{T0} refers to the isothermal bulk modulus. Mao et al. (2011) and Chang et al. (2015) both evaluated the same hydrous Fo_{89} wadsleyite as this study (Z0570) with ~2 wt% water, but Mao et al. (2011) reported $K'_{S0} = 4.8(1)$ using Brillouin spectroscopy, while Chang et al. (2015) obtained $K'_{T0} = 3.77(14)$ from static compression using a BM3-EOS to $P_{\text{max}} = 32$ GPa with ruby pressures and 31 data points. It should be noted, however, that Chang et al. (2015) did not have an experimental V_0 value, which is an important anchor in fitting K'_{T0} with the BM3-EOS because the derivative is effectively evaluated at zero pressure. For comparison, we found $K'_{T0} = 4.26(23)$ for Z0570 in this study using the BM3-EOS to $P_{\text{max}} = 33$ GPa and 22 P - V points. Chang et al. (2015) suggested that the discrepancy in K'_0 between their study and that of Mao et al. (2011) was potentially due to the way density was estimated at high pressure by Mao et al. (2011) by iteration, potentially leading to a higher K'_0 .

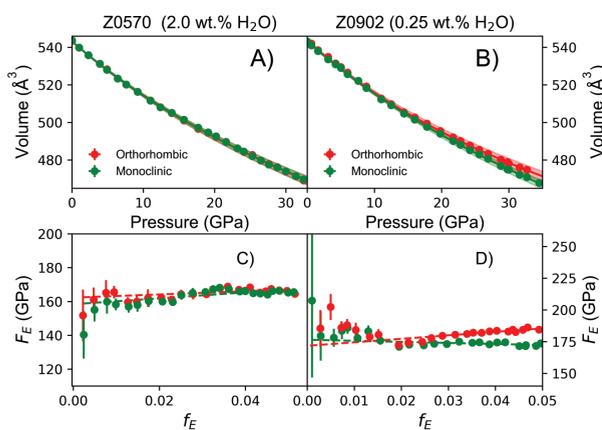


FIGURE 6. Unit-cell volume vs. pressure data of (a) more-hydrous wadsleyite (Z0570, 2.0 wt% H_2O) and (b) less-hydrous wadsleyite (Z0902, 0.25 wt% H_2O) in both orthorhombic (red) and monoclinic symmetries (green) with solid lines showing the fitted third-order BM-EOS. Panels c and d show the corresponding normalized pressure (F_E) vs. Eulerian strain (f_E), where the symmetry is either orthorhombic (red) or monoclinic (green). (Color online.)

We examined the effects of the selection of different crystal symmetry (monoclinic vs. orthorhombic) on the equations of state of wadsleyite by comparing XRD data evaluated using either orthorhombic or monoclinic symmetry. In Figure 6, we compare P - V and F_E - f_E plots of the compression data using orthorhombic symmetry vs. using the observed monoclinic β -angles in calculating volume. The results are compiled in Online Materials¹ Tables S3–S6 and S7. When fitting our P - V data to the BM3-EOS assuming orthorhombic symmetry for the more hydrous wadsleyite (Z0570) we obtain: $V_0 = 543.19(23)$ \AA^3 , $K_{T0} = 164.0(2.7)$ GPa, and $K'_{T0} = 4.26(23)$, whereas on using the monoclinic angles, we obtain: $V_0 = 543.25(23)$ \AA^3 , $K_{T0} = 162.5(2.6)$ GPa, and $K'_{T0} = 4.45(22)$. In short, these EOS parameters agree within their mutual uncertainties, which is not surprising as the β -angle is not very distorted from 90° , and the distortion only showed a quite gradual increase with pressure above 10 GPa for sample Z0570. However, with the less hydrous wadsleyite (Z0902), the BM3-EOS parameters assuming orthorhombic symmetry are: $V_0 = 543.75(29)$ \AA^3 , $K_{T0} = 165.7(3.4)$ GPa, and $K'_{T0} = 5.51(28)$, whereas when using monoclinic symmetry and the observed monoclinic β -angles, we obtain: $V_0 = 543.09(22)$ \AA^3 , $K_{T0} = 172.3(2.3)$ GPa, and $K'_{T0} = 3.91(15)$. Assuming a fixed orthorhombic symmetry at all pressures produced lower K_{T0} values, but the most pronounced effect is the higher first pressure derivative, K'_{T0} .

Hydrogen bonds at high pressure

The long O1-O4 edge of the M3 site is the primary location for hydrogen bonds in wadsleyite, as previously determined by neutron diffraction (Purevjav et al. 2016). While the observed incompressibility of the long O1-O4 edge in the very hydrous sample (Z0570) above 25 GPa might be interpreted as strong hydrogen bonding, the fact that the long O1-O4 edge does not shorten below 2.90(1) \AA , even up to the highest pressure (Online Materials¹ Fig. S5), indicates that the long O1-O4 hydrogen-bonded edge is still far longer than what is considered strong hydrogen bonding, as defined by $d(\text{O}\cdots\text{O}) < 2.5\text{--}2.7$ \AA (Libowitzky 1999). In contrast, high-pressure FTIR data (Yang et al. 2014) up to ~20 GPa and high-pressure Raman data up to ~50 GPa (Kleppe et al. 2006) on hydrous wadsleyite show that the primary O-H stretching modes at 3360–3320 cm^{-1} (at 1 atm) gradually shift to a lower frequency, reaching values below ~3200 cm^{-1} above about 20 GPa, consistent with strong hydrogen bonding (Libowitzky 1999).

The primary hydrogen bond in wadsleyite, therefore, deviates significantly from the relation between O \cdots O distances and O-H stretching frequencies above 20 GPa (Fig. 7). The simplest explanation we can propose, though not prove, involves a shift of the H atom off the longer O1-O4 edge of the M3 site either into the vacant M3 site or more along the c -axis above O1, either way making the O1-H \cdots O4 H bond angle even lower than the observed angle of 171° at room pressure (Purevjav et al. 2016). If the hydrogen bond becomes more nonlinear with pressure, the H atom may move further away from the donor oxygen, and O-H stretching frequency may continue to drop, decoupling the O-H stretching frequency from the relation to hydrogen bond length for linear hydrogen bonds (Libowitzky 1999). Future high-pressure polarized FTIR measurements or high-pressure neutron diffraction experiments of hydrous wadsleyite under hydrostatic compression might be performed to test this hypothesis.

IMPLICATIONS

Wadsleyite is one of the most important minerals that control the physical and chemical properties of the mantle transition zone. To better understand the effects of hydration on the physical properties of wadsleyite at the atomic scale, single-crystal X-ray structure refinements of very hydrous (2.0 wt% H₂O) and slightly hydrous (0.25 wt% H₂O) Fo₉₀ wadsleyite were used to evaluate the crystal structure and hydrogen bond distances up to 32 GPa.

At pressures exceeding 9 GPa, hydrous Fo₉₀ wadsleyite deviates from orthorhombic symmetry to have monoclinic symmetry, which we interpret as resulting from the difference in compression behavior of the split M3 site due to Mg-Fe and vacancy ordering. An increase in the β -angle of hydrous wadsleyite with pressure has implications for the equation of state parameters if lattice parameters are erroneously fitted to orthorhombic symmetry. When the monoclinic symmetry is used in high-pressure EOS fits, our results suggest that K_{70} is lowered while K'_{70} increases with the hydration of wadsleyite. In wadsleyite, hydroxylation is associated with cation vacancies, and vacancies have been found to increase compressibility and K'_{70} of other minerals such as spinels (Nestola et al. 2009). The demonstrated difference in calculated volume due to assumed orthorhombic symmetry, especially at $P > 10$ GPa where the monoclinic angle may increase, could partly explain some aspects of the controversial relationship between water content and K'_{70} . For example, the K'_{70} values obtained by static compression for the same samples in this study (Z0570 and Z0902) and reported in Chang et al. (2015) assuming orthorhombic symmetry were identical and anomalously low with $K'_{70} = 3.7(2)$, whereas we obtained $K'_{70} = 4.5(2)$ for Z0570 and $K'_{70} = 3.9(2)$ for Z0902. Therefore, our results show that ~ 2 wt% H₂O decreases the bulk modulus by about $\sim 6\%$ and increases K'_{70} by about 17%.

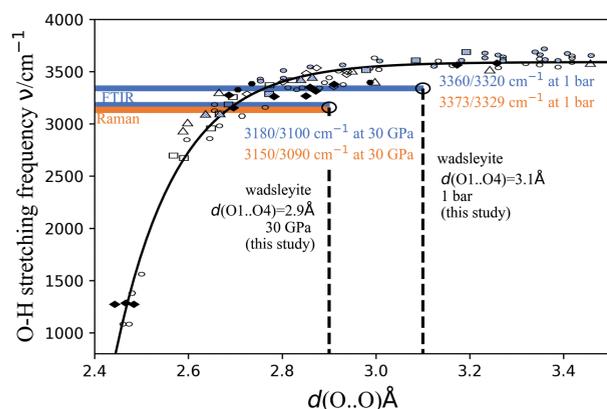


FIGURE 7. Compilation of hydrogen bond distances, $d(\text{O}\cdots\text{O})$, vs. O-H stretching frequencies for nearly linear hydrogen bonds in minerals, modified from Libowitzky (1999). Large circles indicate the intersection of measured $d(\text{O}\cdots\text{O})$ for the O1-H \cdots O4 hydrogen bond in wadsleyite from this study with measured spectroscopic O-H stretching frequencies from FTIR (Yang et al. 2014) and Raman studies (Kleppe et al. 2006). The FTIR O-H stretching frequencies of 3360 and 3320 cm^{-1} at 1 bar are from Yang et al. (2014), and at 30 GPa (blue band), they are calculated from the polynomial fits published by Yang et al. (2014) from data up to 20 GPa. The Raman O-H stretching frequencies at 30 GPa of 3090 and 3150 cm^{-1} are from the study of Kleppe et al. (2006). (Color online.)

The decrease of K_{70} and increase of K'_{70} can cause a volume crossover on compression (Fig. 2), although the more hydrous sample remains less dense at high pressure.

The primary hydrogen bond in wadsleyite, defined by O1-H \cdots O4 along the longer O1-O4 edge of the M3 site, becomes incompressible at pressures above 25 GPa. Although strong hydrogen bonding in wadsleyite above ~ 15 GPa is implied from previous spectroscopic studies showing O-H stretching frequencies that dropped below 3200 cm^{-1} (Yang et al. 2014; Kleppe et al. 2006), we found that the hydrogen bond distance, $d(\text{O1}\cdots\text{O4})$, does not drop below 2.9 Å in length and certainly not into the region of strong hydrogen bonding at 2.5–2.7 Å. We conclude that the primary hydrogen bond in wadsleyite may become more nonlinear at high pressures. This is the simplest explanation for the observed combination of a long O \cdots O distance and low O-H stretching frequency and could be explained if the O atom moves off the long O1-O4 edge into the vacant M3 site or out above O1 more along the c -axis. Due to isotopic effects on hydrogen-bond potential energy functions (e.g., Novak 1974), short, strong hydrogen bonds are expected to be enriched in hydrogen vs. deuterium (e.g., Kuroda et al. 1979). Even though the hydrogen-bond length in wadsleyite determined here at transition zone pressures would not predict strong isotopic fractionation (δD), the observed O-H stretching frequencies below 3300 cm^{-1} (Yang et al. 2014) indicate otherwise, highlighting the importance of combining structure (XRD) and spectroscopic evidence (e.g., Fig. 7) to evaluate hydrogen-bonding geometry at pressures where neutron diffraction is not yet feasible.

DATA AVAILABILITY STATEMENT

All structure data as CIF have been uploaded to the American Mineralogist Crystal Structure Database.

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REFERENCES CITED

- Angel, R.J., Alvaro, M., and Gonzalez-Platas, J. (2014) Eosfit7c and a Fortran module (library) for equation of state calculations. *Zeitschrift für Kristallographie: Crystal-Line Materials*, 229, 405–419, <https://doi.org/10.1515/zkri-2013-1711>.
- Buchen, J., Marquardt, H., Boffa Ballaran, T., Kawazoe, T., and McCammon, C. (2017) The equation of state of wadsleyite solid solutions: Constraining the effects of anisotropy and crystal chemistry. *American Mineralogist*, 102, 2494–2504, <https://doi.org/10.2138/am-2017-6162>.
- Buchen, J., Marquardt, H., Speziale, S., Kawazoe, T., Boffa Ballaran, T., and Kumosov, A. (2018) High-pressure single-crystal elasticity of wadsleyite and the seismic signature of water in the shallow transition zone. *Earth and Planetary Science Letters*, 498, 77–87, <https://doi.org/10.1016/j.epsl.2018.06.027>.

- Chang, Y.-Y., Jacobsen, S.D., Bina, C.R., Thomas, S.-M., Smyth, J.R., Frost, D.J., Boffa Ballaran, T., McCammon, C.A., Hauri, E.H., Inoue, T., and others. (2015) Comparative compressibility of hydrous wadsleyite and ringwoodite: Effect of H₂O and implications for detecting water in the transition zone. *Journal of Geophysical Research: Solid Earth*, 120, 8259–8280, <https://doi.org/10.1002/2015JB012123>.
- Chopelas, A. (1991) Thermal properties of β -Mg₂SiO₄ at mantle pressures derived from vibrational spectroscopy: Implications for the mantle at 400 km depth. *Journal of Geophysical Research*, 96 (B7), 11817–11829, <https://doi.org/10.1029/91JB00898>.
- Cynn, H. and Hofmeister, A.M. (1994) High pressure IR spectra of lattice modes and OH vibrations in Fe-bearing wadsleyite. *Journal of Geophysical Research*, 99 (B9), 17717–17727, <https://doi.org/10.1029/94JB01661>.
- Deon, F., Koch-Muller, M., Rhede, D., Gottschalk, M., Wirth, R., and Thomas, S.-M. (2010) Location and quantification of hydroxyl in wadsleyite: New insights. *American Mineralogist*, 95, 312–322, <https://doi.org/10.2138/am.2010.3267>.
- Dolomanov, O.V., Bourhis, L.J., Gildea, R.J., Howard, J.A., and Puschmann, H. (2009) Olex2: A complete structure solution, refinement and analysis program. *Journal of Applied Crystallography*, 42, 339–341, <https://doi.org/10.1107/S0021889808042726>.
- Dong, J.J., Fischer, R.A., Stixrude, L.P., and Lithgow-Bertelloni, C.R. (2021) Constraining the volume of Earth's early oceans with a temperature-dependent mantle water storage capacity model. *AGU Advances*, 2(1), e2020AV000323.
- Emsley, J. (1980) Very strong hydrogen bonding. *Chemical Society Reviews*, 9, 91–124, <https://doi.org/10.1039/c9800900091>.
- Fei, H. and Katsura, T. (2020) High water solubility of ringwoodite at mantle transition zone temperature. *Earth and Planetary Science Letters*, 531, 115987, <https://doi.org/10.1016/j.epsl.2019.115987>.
- Gwanmesia, G.D., Whitaker, M.L., Dai, L., James, A., Chen, H., Triplett, R.S., and Cai, N. (2020) The elastic properties of β -Mg₂SiO₄ containing 0.73 wt% of H₂O to 10 GPa and 600 K by ultrasonic interferometry with synchrotron X-radiation. *Minerals*, 10, 209, <https://doi.org/10.3390/min10030209>.
- Hazen, R.M. (1993) Comparative compressibilities of silicate spinels: anomalous behavior of (Mg,Fe)₂SiO₄. *Science*, 259(5092), 206–209, <https://doi.org/10.1126/science.259.5092.206>.
- Hazen, R.M., Weinberger, M.B., Yang, H., and Prewitt, C.T. (2000) Comparative high-pressure crystal chemistry of wadsleyite, β -(Mg_{1-x}Fe_x)₂SiO₄, with x=0 and 0.25. *American Mineralogist*, 85, 770–777, <https://doi.org/10.2138/am-2000-5-617>.
- Holl, C.M., Smyth, J.R., Jacobsen, S.D., and Frost, D.J. (2008) Effects of hydration on the structure and compressibility of wadsleyite, β -Mg₂SiO₄. *American Mineralogist*, 93, 598–607, <https://doi.org/10.2138/am.2008.2620>.
- Horiuchi, H. and Sawamoto, H. (1981) β -Mg₂SiO₄: Single-crystal X-ray diffraction study. *American Mineralogist*, 66, 568–575.
- Jacobsen, S.D., Demouchy, S., Frost, D.J., Boffa Ballaran, T., and Kung, J. (2005) A systematic study of OH in hydrous wadsleyite from polarized FTIR spectroscopy and single-crystal X-ray diffraction: Oxygen storage in hydrogen storage in Earth's interior. *American Mineralogist*, 90, 61–70, <https://doi.org/10.2138/am.2005.1624>.
- Jacobsen, S.D., Holl, C.M., Adams, K.A., Fischer, R.A., Martin, E.S., Bina, C.R., Lin, J.-F., Prakapenka, V.B., Kubo, A., and Dera, P. (2008) Compression of single-crystal magnesium oxide to 118 GPa and a ruby pressure gauge for helium pressure media. *American Mineralogist*, 93, 1823–1828, <https://doi.org/10.2138/am.2008.2988>.
- Karato, S.I. (2011) Water distribution across the mantle transition zone and its implications for global material circulation. *Earth and Planetary Science Letters*, 301, 413–423, <https://doi.org/10.1016/j.epsl.2010.11.038>.
- Kleppe, A., Jephcoat, A., Olijnyk, H., Slesinger, A., Kohn, S., and Wood, B. (2001) Raman spectroscopic study of hydrous wadsleyite (β -Mg₂SiO₄) to 50 GPa. *Physics and Chemistry of Minerals*, 28, 232–241, <https://doi.org/10.1007/s002690100152>.
- Kleppe, A.K., Jephcoat, A.P., and Smyth, J.R. (2006) High-pressure Raman spectroscopic study of Fe₃₀ hydrous wadsleyite. *Physics and Chemistry of Minerals*, 32, 700–709, <https://doi.org/10.1007/s00269-005-0048-8>.
- Kohn, S., Brooker, R.A., Frost, D.J., Slesinger, A., and Wood, B. (2002) Ordering of hydroxyl defects in hydrous wadsleyite (β -Mg₂SiO₄). *American Mineralogist*, 87, 293–301, <https://doi.org/10.2138/am-2002-2-310>.
- Kuroda, Y., Suzuoki, T., and Matsuo, S. (1979) The lowest δ D value found in a hydrous silicate, pectolite. *Nature*, 279, 227–228, <https://doi.org/10.1038/279227a0>.
- Li, B. and Liebermann, R.C. (2000) Sound velocities of wadsleyite β -(Mg_{0.88}Fe_{0.12})₂SiO₄ to 10 GPa. *American Mineralogist*, 85, 292–295, <https://doi.org/10.2138/am-2000-2-305>.
- Libowitzky, E. (1999) Correlation of OH stretching frequencies and OH–O hydrogen bond lengths in minerals. *Monatshefte für Chemie/Chemical Monthly*, 130(8), 1047–1059, <https://doi.org/10.1007/BF03354882>.
- Liu, W., Kung, J., Li, B., Nishiyama, N., and Wang, Y. (2009) Elasticity of (Mg_{0.8}Fe_{0.13})₂SiO₄ wadsleyite to 12 GPa and 1073 K. *Physics of the Earth and Planetary Interiors*, 174, 98–104, <https://doi.org/10.1016/j.pepi.2008.10.020>.
- Mao, Z., Jacobsen, S.D., Frost, D.J., McCammon, C.A., Hauri, E.H., and Duffy, T.S. (2011) Effect of hydration on the single-crystal elasticity of Fe-bearing wadsleyite to 12 GPa. *American Mineralogist*, 96, 1606–1612, <https://doi.org/10.2138/am.2011.3807>.
- McMillan, P.F., Akaogi, M., Sato, R.K., Poe, B., and Foley, J. (1991) Hydroxyl groups in β -Mg₂SiO₄. *American Mineralogist*, 76, 354–360.
- Momma, K. and Izumi, F. (2011) VESTA 3 for three-dimensional visualization of crystal, volumetric and morphology data. *Journal of Applied Crystallography*, 44, 1272–1276, <https://doi.org/10.1107/S0021889811038970>.
- Nestola, F., Smyth, J.R., Parisatto, M., Secco, L., Princivalle, F., Bruno, M., Prencipe, M., and Dal Negro, A. (2009) Effects of non-stoichiometry on the spinel structure at high pressure: Implications for Earth's mantle mineralogy. *Geochimica et Cosmochimica Acta*, 73, 489–492, <https://doi.org/10.1016/j.gca.2008.11.001>.
- Novak, A. (1974) Hydrogen bonding in solids correlation of spectroscopic and crystallographic data. In J.-H. Fuhrhop, G. Blauer, T.J.R. Weakley, and A. Novak, Eds., *Large Molecules*, p. 177–216. Springer, <https://doi.org/10.1007/BFb0116438>.
- Pearson, D.G., Brenker, F.E., Nestola, F., McNeill, J., Nasdala, L., Hutchison, M.T., Matveev, S., Mather, K., Silversmit, G., Schmitz, S., and others. (2014) Hydrous mantle transition zone indicated by ringwoodite included within diamond. *Nature*, 507, 221–224, <https://doi.org/10.1038/nature13080>.
- Purejav, N., Okuchi, T., Tomioka, N., Wang, X., and Hoffmann, C. (2016) Quantitative analysis of hydrogen sites and occupancy in deep mantle hydrous wadsleyite using single crystal neutron diffraction. *Scientific Reports*, 6, 34988, <https://doi.org/10.1038/srep34988>.
- Rivers, M., Prakapenka, V.B., Kubo, A., Pullins, C., Holl, C.M., and Jacobsen, S.D. (2008) The COMPRES/GSECARS gas-loading system for diamond anvil cells at the Advanced Photon Source. *High Pressure Research*, 28, 273–292, <https://doi.org/10.1080/08957950802333593>.
- Sano-Furukawa, A., Kuribayashi, T., Komatsu, K., Yagi, T., and Ohtani, E. (2011) Investigation of hydrogen sites of wadsleyite: A neutron diffraction study. *Physics of the Earth and Planetary Interiors*, 189, 56–62, <https://doi.org/10.1016/j.pepi.2011.07.003>.
- Sheldrick, G.M. (2015) Crystal structure refinement with SHELXL. *Acta Crystallographica*, C71, 3–8, <https://doi.org/10.1107/S205229614024218>.
- Smyth, J.R. (1987) The β -Mg₂SiO₄: a potential host for water in the mantle? *American Mineralogist*, 72, 1051–1055.
- (1994) A crystallographic model for hydrous wadsleyite (β -Mg₂SiO₄): An ocean in the Earth's interior? *American Mineralogist*, 79, 1021–1024.
- Smyth, J.R., Kawamoto, T., Jacobsen, S.D., Swope, R.J., Hervig, R.L., and Holloway, J.R. (1997) Crystal structure of monoclinic hydrous wadsleyite [β -(Mg,Fe)₂SiO₄]. *American Mineralogist*, 82, 270–275, <https://doi.org/10.2138/am-1997-3-404>.
- Smyth, J.R., Bolfan-Casanova, N., Avignone, D., El-Ghozzi, M., and Himer, S.M. (2014) Tetrahedral ferric iron in oxidized hydrous wadsleyite. *American Mineralogist*, 99, 458–466, <https://doi.org/10.2138/am.2014.4520>.
- Sturhahn, W. (2000) CONUSS and PHOENIX: Evaluation of nuclear resonant scattering data. *Hyperfine Interactions*, 125, 149–172, <https://doi.org/10.1023/A:1012681503686>.
- Toellner, T.S. (2000) Monochromatization of synchrotron radiation for nuclear resonant scattering experiments. *Hyperfine Interactions*, 125, 3–28, <https://doi.org/10.1023/A:1012621317798>.
- Tsuchiya, J. and Tsuchiya, T. (2009) First principles investigation of the structural and elastic properties of hydrous wadsleyite under pressure. *Journal of Geophysical Research*, 114 (B2), B02206, <https://doi.org/10.1029/2008JB005841>.
- Wang, F., Barklage, M., Lou, X., van der Lee, S., Bina, C.R., and Jacobsen, S.D. (2018) HyMaTZ: A Python program for modeling seismic velocities in hydrous regions of the mantle transition zone. *Geochemistry, Geophysics, Geosystems*, 19, 2308–2324, <https://doi.org/10.1029/2018GC007464>.
- Yang, X., Keppler, H., Dubrovinsky, L., and Kurnosov, A. (2014) In-situ infrared spectra of hydroxyl in wadsleyite and ringwoodite at high pressure and high temperature. *American Mineralogist*, 99, 724–729, <https://doi.org/10.2138/am.2014.4634>.
- Ye, Y., Smyth, J.R., Hushur, A., Manghnani, M.H., Lonappan, D., Dera, P., and Frost, D.J. (2010) Crystal structure of hydrous wadsleyite with 2.8% H₂O and compressibility to 60 GPa. *American Mineralogist*, 95, 1765–1772, <https://doi.org/10.2138/am.2010.3533>.
- Ye, Y., Smyth, J.R., and Frost, D.J. (2011) Structural study of the coherent dehydration of wadsleyite. *American Mineralogist*, 96, 1760–1767, <https://doi.org/10.2138/am.2011.3852>.
- Zhang, D., Dera, P.K., Eng, P.J., Stubbs, J.E., Zhang, J.S., Prakapenka, V.B., and Rivers, M.L. (2017) High pressure single crystal diffraction at PX². *Journal of Visualized Experiments*, 119, 54660, <https://doi.org/10.3791/54660>.

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