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Key Points:

- ε-type iron carbonitrides undergo two phase transitions at high pressure and high temperature
- N/C and Fe/(N+C) atomic ratios are two key factors affecting the phase stability and elasticity of iron carbonitrides
- Iron carbonitrides with diverse structures may be the main host for nitrogen in the deep mantle and encapsulated in superdeep diamonds

Supporting Information:

Supporting Information may be found in the online version of this article.

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High-Pressure Phase Stability and Thermoelastic Properties of Iron Carbonitrides and Nitrogen in the Deep Earth

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Abstract Iron-dominant metallic phases are likely the primary hosts for nitrogen in the reduced deep Earth, hence the storage of nitrogen in the lower mantle and the core is governed by the behavior of the Fe-N-C system at high temperatures and pressures. In this study, phase transitions and thermoelastic properties of iron carbonitrides were investigated at high pressure-temperature conditions by diamond anvil cell experiments and first-principles calculations. Experimental data revealed no phase transition in ε -type Fe₄ (N_{0.6}C_{0.4}) or Fe₇ (N_{0.75}C_{0.25})₃ up to 60 GPa at room temperature. At high temperature, Fe₇ $(N_{0.75}C_{0.25})_3$ transforms into the Fe₃C-type phase at ~27 GPa, and then into the Fe₇C₃-type phase at \sim 45 GPa, which is also corroborated by our theoretical calculations. We found that the phase stability of iron carbonitrides mainly depends on the N/C ratio, and the elastic properties of iron carbonitrides are dominantly affected by the Fe/(N+C) ratio. Iron carbonitrides with diverse structures may be the main host for nitrogen in the deep mantle. Some iron carbonitride inclusions in lower mantle diamonds could be the residue of the primordial mantle or originate from subducted nitrogen-bearing materials, rather than iron-enriched phases of the outer core. In addition, our experiments confirmed the existence of Fe_7C_3 -type Fe_7C_3 - Fe_7N_3 solid solutions above 40 GPa. Fe_7C_3 -type $Fe_7(C, N)_3$ has comparable density and thermoelastic properties to its isostructural endmembers and may be a promising candidate constituent of the Earth's inner core.

Plain Language Summary Nitrogen is an essential element for the Earth's atmosphere and life. Most of the Earth's nitrogen may reside in Earth's interior, which may profoundly influence the partial pressure of atmospheric nitrogen. Knowledge on the storage and cycling of nitrogen in the deep Earth is crucial for our understanding of the Earth's evolution and dynamics. We combined high-pressure and high-temperature experiments and theoretical calculations to investigate the phase transitions and elasticity of iron carbonitrides. The results show that iron carbonitrides are stable upon compression at room temperature, but transformed into two high-pressure structures at lower pressures at high temperatures than their iron carbide counterparts. N/C and Fe/(N+C) atomic ratios are two key factors affecting the phase stability and elasticity of iron carbonitrides. Nitrogen could be stored in the Earth's lower mantle in the form of iron carbonitrides. Iron carbonitrides may be the main nitrogen host and be trapped as inclusions in some superdeep diamonds and carried to the shallow regions through geodynamic processes, as we can find in some diamonds of sublithospheric origin.

1. Introduction

Nitrogen (N) is one of the most significant volatiles in our planet, an essential element for life, and the primary component of the atmosphere. Nitrogen in the Earth's interior is linked to that near the surface through processes of subduction and volcanic outgassing, influencing the partial pressure (*P*) of atmospheric N in a profound way (Busigny et al., 2019; Mikhail & Sverjensky, 2014). This "deep nitrogen" also provides crucial insights into planetary evolution, including the accretion and differentiation of the core and mantle

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on Earth, as well as the delivery of late veneers and the deep recycling of other volatiles associated with N (Bergin et al., 2015; Dalou et al., 2017; Grewal, Dasgupta, Sun, et al., 2019; Kaminsky & Wirth, 2017). Based upon the abundance of N in the bulk Earth relative to that measured in carbonaceous chondrites, the bulk Earth is depleted by an order of magnitude in N in comparison with other volatiles, which is the so-called "missing nitrogen" problem (Marty, 2012). A recent study by Johnson and Goldblatt (2015) has suggested that N contents in the "deep Earth" may be considerably underestimated, and that more than 90% of the bulk Earth N content is likely to be residing in these hidden and largely "inaccessible" reservoirs. Core-mantle differentiation represents a first-order process controlling the initial distribution of N among the various reservoirs in the deep Earth. The partitioning behavior of N between iron-enriched alloys and silicates can be affected by P, temperature (T), and the composition of the participating phases, but the prevailing oxygen fugacity (f_{02}) appears to be the dominant factor (Dalou et al., 2017; Grewal, Dasgupta, Holmes, et al., 2019; Grewal, Dasgupta, Sun, et al., 2019; Roskosz et al., 2013). At realistic *P*-*T*- f_{O_2} conditions for core formation (i.e., 40~60 GPa, 3,000~4000 K and IW-3~0), N prefers to bond with iron (Fe), and is incorporated more into the Earth's core than into the silicate mantle. On the other hand, the silicate mantle may be saturated with Fe at depths >250 km, and Fe concentration can reach ~1 wt% in the lower mantle due to disportionation of Fe²⁺ and preferential incorporation of Fe³⁺ in bridgmanite (Frost et al., 2004). The partition coefficient for N between metallic Fe and bridgmanite is close to 100, indicating that the N content of the metallic phase in the lower mantle should be nearly equivalent to that of the rest of the silicate lower mantle (Yoshioka et al., 2018). Therefore, the Fe-enriched metallic phase is likely to be a main host for N in the Earth's interior. A more complete picture of the N evolution of the Earth requires the determination of the composition and stability of N-bearing Fe-enriched minerals at high P-T conditions.

In the Fe-N system, several Fe nitrides are stable at ambient conditions, including cubic γ' -Fe₄N (space group: $Pm\overline{3}m$), orthorhombic ζ -Fe₂N (space group: Pbcn) and nonstoichiometric ε -type Fe₃N_x (Wriedt et al., 1987). The space group of ε -type Fe₃N_x has been assigned to either $P3_12$ or $P6_322$, in which N atoms occupy part of the octahedral vacancies. The ε -type Fe₃N_x is capable of holding a range of N in its structure, with x varying from 0.6 to 1.5 (Figure S1) (Litasov et al., 2017; Niewa, Rau, Wosylus, Meier, Hanfland, et al., 2009). At 9~15 GPa and 1,400~1600 K, Fe₄N and Fe₂N undergo structural transitions to ε -type phases (Guo et al., 2013; Niewa, Rau, Wosylus, Meier, Wessel, et al., 2009; Schwarz et al., 2009). The ε -type Fe₃N_x is stable up to ~40 GPa and ~1500 K (Litasov et al., 2017; Minobe et al., 2015). Upon further compression, ε -type Fe₃N_x transforms into Fe₇C₃-type Fe₇N₃ (Minobe et al., 2015). As compared to the ε -type phase, Fe₇C₃-type Fe₇N₃ is stable over a much broader *P*-*T* range that is, to at least ~150 GPa and ~3,000 K (Kusakabe et al., 2019; Minobe et al., 2015). These results suggest that, among various Fe nitrides, ε -type Fe₃N_x and Fe₇C₃-type Fe₇N₃ are strong candidates for the storage hosts of N in the Earth's mantle and core.

Because of relatively low abundance of N in the bulk Earth, it is likely that, in addition to forming pure Fe nitrides, N substitutes for other light elements into the crystal structures of Fe-enriched alloys. Carbon (C) has similar geochemical and cosmochemical properties to N. They may mutually coexist and replace each other to form solid solutions in deep reservoirs (Sagatov et al., 2019). Indeed, recent studies have reported various Fe carbonitrides and nitrocarbides, such as $Fe_9(N_{0.8}C_{0.2})_4$, $Fe_5(C_{0.48}N_{0.52})_2$ and $Fe_7(C_{0.73}N_{0.27})_3$, inclusions in superdeep diamonds (Kaminsky & Wirth, 2017). A recent study shows that N could replace C in Fe₃C to form an ε -type C-bearing Fe₃N_x at 7~17 GPa, indicating that N appears to have a better affinity to Fe than C at high P-T conditions (Litasov et al., 2016). The combined effects of N and C in the Fe-enriched Fe-N-C system under simultaneous high-P and high-T conditions provide indispensable clues to decipher the storage and speciation of N in the Earth's mantle and core. Sokol et al. (2017) have recently reported a complex phase diagram for the Fe-enriched Fe-N-C system at 7.8 GPa and 1,350°C. However, there are no experimental or theoretical studies to examine the phase stability and physical properties of Fe-N-C compounds at high P-T conditions. In this paper, we report the phase transitions and thermoelastic properties of two synthetic Fe carbonitrides of different stoichiometry at extreme conditions up to 60 GPa and 2,000 K, as determined in experiments carried out using externally heated and laser-heated diamond anvil cells (EH-DAC and LHDAC) combined with synchrotron-based X-ray diffraction (XRD). In addition, first-principles calculations based on density functional theory (DFT) were performed to complement our experimental observations. Our results not only contribute to an improved understanding of the speciations of N stored in the Earth's deep interior but also shed light on the formation mechanism of Fe-N-C inclusions in lower mantle diamonds.



2. Materials and Methods

2.1. Sample Synthesis and Characterization

Two ε -type Fe carbonitrides were synthesized at high *P*-*T* conditions using a 2000-ton multi-anvil press at the University of Hawai'i at Mānoa. Starting materials were the mixtures of Fe₃N powder and synthetic Fe₃C. The Fe₃N powder was commercially obtained from Kojundo Chemical Lab. Co. Ltd. The Fe₃C sample was synthesized from a mixture of Fe powder (99.9+% purity, Aldrich Chemical Company) and graphite powder (99.9995% purity, Alfa Aesar Company) at an atomic ratio of Fe:C = 3:1 at 3 GPa and 1,300 K for 8 h using a MgO capsule and a graphite heater. The first ε -type Fe carbonitride sample was synthesized at 8 GPa and 1,500 K for 1.5 h using a MgO capsule and a 14/8 cell assembly (Liu et al., 2020). The synthesis of the second ε -type Fe carbonitride sample was carried out at 9 GPa and 1,400 K for 20 h using a MgO capsule and a 10/5 cell assembly (Leinenweber et al., 2012).

Back-scattered electron (BSE) images of individual samples were collected using the scanning electron microscope (SEM) on Cameca SX100 at the Robert B. Mitchell Electron Microbeam Laboratory (RM-EML), Department of Earth and Environmental Sciences, University of Michigan, Ann Arbor, and School of Earth and Space Sciences, Peking University. BSE images show that each synthetic sample consists of a homogeneous phase, and in addition to Fe carbonitride product, there is little graphite in the first sample (Figure S2).

Concentrations of Fe, N and C were determined by Cameca SX100 electron probe microanalysis (EPMA) at the RM-EML, using a procedure similar to that of von der Handt and Dalou (2016) and Dalou et al. (2017). The accelerating potential was 10 kV, the beam current was 200 nA and the beam spot was focused to 1 μ m. Fe Kα X-rays were measured using a large LiF diffracting crystal, using synthetic Fe₃C as a calibration standard; N K α X-rays were measured using a large PC0 pseudocrystal (W/Si multilayer, 2d \approx 4.5 nm), using synthetic Si₃N₄ as a calibration standard; and C Kα X-rays were measured using a PC2 diffracting crystal (Ni/C multilayer, $2d \approx 10$ nm), using synthetic Fe₃C as a calibration standard. Wavelength scans were performed around the Fe K α , N K α , and C K α peaks to accurately determine the off-peak spectrometer positions for measuring continuum, and for modeling the continuum curvature (von der Handt & Dalou, 2016). For N $K\alpha$ measurements, widely spaced off-peak positions were used when measuring the Si₃N₄ standard due to the broadness of the N K α peak from Si₃N₄. Off-peak measurements were made significantly closer to the peak on Fe-N-C samples, in order to reduce the uncertainty of the interpolation and to avoid the interference of the nearby second Bragg order Fe L α and L β lines with the off-peak measurements (Figure S3); these intensities of these lines are significantly reduced by the PC0 pseudocrystal, but they are not fully suppressed even when differential pulse-height analyzer mode is used. Continuum intensities at the onpeak positions were interpolated from off-peak continuum measurements using an exponential function for N K α and C K α and a linear function for Fe K α (von der Handt & Dalou, 2016). Counting times for all measured X-ray lines were 20 s on the peak and 10 s on each off-peak position. Samples were coated with \sim 15 nm of aluminum prior to analysis. In order to minimize the effects of differential coating thicknesses, calibration standards and Fe-N-C samples used in this study were coated at the same time. A liquid nitrogen cold finger was used to minimize carbon contamination of the sample surface during analysis, and a small time-dependent intensity correction was applied to N Ka and C Ka count rate data to account for changes in X-ray intensity due to degradation of the sample or conductive coating during analysis; no systematic changes in Fe K α count rates over the course of the analyses were observed. Concentrations were calculated from measured, continuum-corrected X-ray intensities using the PROZA matrix corrections and the FFAST mass absorption coefficients (von der Handt & Dalou, 2016). In addition, an interference correction was applied to the measured N Ka intensities, to account for intensities contributed from the second Bragg order Fe L β 3 and L β 4 lines near the N K α peak (Figure S3); the magnitude of the correction was <5%. Suitable secondary standards containing Fe, N and C were not available to verify accuracy. In lieu of such materials, a synthetic, nominally nitrogen-free Fe-C sample was measured to verify that N concentrations were at or near zero. The N concentration measured in this material was $10 \pm 250 \,\mu$ g/g. Also, any measurements with analytical totals of less than 98.5% or greater than 101.5% were excluded.

EPMA results show that the two samples are chemically homogeneous, with the average composition of the first sample as $Fe_4(N_{0.6}C_{0.4})$, and the second sample as $Fe_7(N_{0.75}C_{0.25})_3$ (Table S1). The crystal structures of these two samples were determined by XRD at beamline 13-BMC, Advanced Photon Source (APS), Argonne



National Laboratory (ANL). The XRD measurements confirm that both the $Fe_4(N_{0.6}C_{0.4})$ and $Fe_7(N_{0.75}C_{0.25})_3$ phases adopt the ε -type structure.

2.2. High P-T Synchrotron XRD Experiments

BX90-type DACs with diamond anvils of 250 μ m (in run-1, 2 and 4) or 300 μ m (in run-3) were used to generate high *P* in four experimental runs. 130~160 μ m diameter holes were drilled in pre-indented 30~35 μ m rhenium gaskets to serve as sample chambers. Fe₄(N_{0.6}C_{0.4}), with a thickness of ~15 μ m, was loaded into the sample chamber for run-1, with Fe₇(N_{0.75}C_{0.25})₃ comprising runs 2 and 3. A piece of Au foil placed near the sample was used to determine the internal *P* in all three of these runs (Fei et al., 2007). Neon gas was loaded as the *P*-transmitting medium (PTM) and ruby spheres were also loaded into the sample chamber to measure the *P* for the gas loading. In run-3, three Pt-10 wt% Rh (0.01") wires wound around a pyrophyllite ring base allowed us to heat the sample to a maximum *T* of 750 K, and two K-type thermocouples (Chromega-Alomega 0.005"" and Chromega-Alomega 0.01"') in contact with the "downstream" diamond were used to measure the *T* (Lai et al., 2020). The difference in measured temperatures between these two thermocouple junctions was smaller than 15 K. In run-4, Fe₇(N_{0.75}C_{0.25})₃ with a thickness of ~10 μ m was sandwiched between two KCl plates. In this arrangement, KCl acts as a PTM, a thermal insulator, and as a *P* indicator (Tateno et al., 2019).

Synchrotron XRD experiments of runs 1–3 were conducted at beamline 13-BMC, APS, ANL. The incident X-ray beam with a wavelength of 0.434 Å was focused to an area approximately 15 μ m in diameter measured at the full width at half maximum. Membrane *P* control system was used to increase *P* inside the DAC automatically. For the EHDAC experiment (run-3), XRD spectra were collected at four different temperatures (300 K, 450 K, 600 K and 750 K) at each given *P*. The LHDAC experiment (run-4) was performed at beamline 13-IDD, APS, ANL. The incident X-ray beam ($\lambda = 0.3344$ Å) had a focal spot size of $3 \times 4 \mu m^2$. The sample was heated gradually from both sides at intervals of 100~200 K at a given *P* using a double-sided laser heating system with the flat-top heating spot size of ~10 μ m in diameter. A fresh spot of the sample was used for laser heating in each heating cycle. The *T* was determined by fitting the thermal radiation data using the Planck radiation function under the gray body approximation. The *P* at high *T* was determined from the volume of KCl using the thermal equation of state (EoS) (Tateno et al., 2019). The effective *T* of KCl was estimated according to Campbell et al. (2009) using the following approximation:

$$T_{\rm KCl} = \frac{3 \times T_{\rm meas} + 300}{4} \pm \frac{T_{\rm meas} - 300}{2}$$

The powder XRD spectra were integrated and then processed using model-biased Le Bail fitting with GSAS + EXPGUI software. For the single-crystal XRD spectra, the *d*-spacings of the sample at various *P*-*T* conditions were extracted using the GSE_ADA software, and lattice parameters were refined by the least squares method. The *P*-volume (*V*) data were fitted using EoSFit7c software (Angel et al., 2014).

2.3. DFT-Based First-Principles Calculations

DFT-based first-principles calculations were performed using the "projected augmented wave (PAW) method" as implemented in Vienna *ab-initio* simulation package (Kresse & Joubert, 1999). The exchange correlation potential was treated by the generalized gradient approximation (GGA) in the Perdew-Burke-Ernzerhof (PBE) version (Perdew et al., 1996). We chose the kinetic energy cut-off as 1,000 eV in static simulations. The energy convergence criterion was 10^{-6} eV for electronic self-consistent calculations. The *k*-points grids were set as $7 \times 7 \times 7$, $6 \times 6 \times 8$ and $5 \times 5 \times 7$ in the Brillouin zone for the trigonal ε -type, orthorhombic Fe₃Ctype and hexagonal Fe₇C₃-type structures, respectively. A $1 \times 1 \times 7$ supercell was constructed to simulate ε -type Fe₇N₃ and Fe₇(N_{5/6}C_{1/6})₃ phases. And a $1 \times 1 \times 2$ supercell was for ε -type Fe₃(N_{0.75}C_{0.25}) and *h*-type Fe₇(N_{0.75}C_{0.25})₃ phases. The total energy difference was converged to 10^{-5} eV/formula unit (f.u.) with respect to the energy cut-off or *k*-points. The corresponding force difference was converged to 10^{-3} eV/Å (less than 0.1 GPa). The spin-polarization of Fe without spin-orbit coupling was considered in the simulations. The ferromagnetic (FM) or nonmagnetic (NM) states were calculated for selected candidates. For each candidate phase, the cell parameters, atomic coordinates, and sublattice magnetic moments were allowed



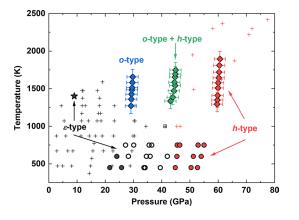


Figure 1. Phase diagram of Fe₇(N_{0.75}C_{0.25})₃ at high *P*-*T* conditions. The ε -type, *o*-type and *h*-type represent the *P*3₁2, Fe₃C-type (*Pnma*) and Fe₇C₃-type (*P6*₃*mc*) phases, respectively. The black star represents the synthetic conditions of ε -type Fe₇(N_{0.75}C_{0.25})₃ (9 GPa and 1,400 K). The open circles represent the metastable ε -type phase at 27~40 GPa and 450~750 K. The open gray and red crosses represent the ε -type and *h*-type phases in the Fe-N binary system, respectively (Litasov et al., 2017; Minobe et al., 2015).

to relax at each given *V*. The calculated *V*-energy (*E*) data were fitted by the Birch-Murnaghan EoS to obtain the *P*-enthalpy (*H*) relation for each phase (Birch, 1947; Murnaghan, 1944):

$$E(V) = E_0 + \frac{9V_0K_0}{16} \left\{ \left[\left(\frac{V_0}{V}\right)^{\frac{2}{3}} - 1 \right]^3 K'_0 + \left[\left(\frac{V_0}{V}\right)^{\frac{2}{3}} - 1 \right]^2 \left[6 - 4\left(\frac{V_0}{V}\right)^{\frac{2}{3}} \right] \right\}$$

where E_0 , V_0 , and K_0 are the energy, unit-cell volume and isothermal bulk modulus at zero *P*, respectively. K_0' is the first *P* derivative of K_0 . The derived *H* of each phase was compared with all the others to ascertain which phase is the most energetically favored at a given *P* and 0 K.

3. Results

3.1. Stability and Phase Transitions of Iron Carbonitrides

Four runs of experiments were conducted to investigate the stability of Fe carbonitrides at high *P*-*T* conditions up to ~60 GPa and ~2000 K (Figure 1 and Table S2). No phase transition was observed at ambient *T* up to ~60 GPa in ε -type Fe₄(N_{0.6}C_{0.4}) (run-1) or to ~50 GPa in ε -type

Fe₇(N_{0.75}C_{0.25})₃ (run-2) (Figure S4). This indicates that the initial ε -type structure is stable over the studied *P* range at ambient *T*.

In run-3, ε -type Fe₇(N_{0.75}C_{0.25})₃ was observed to be stable to 42.0 GPa and 600 K (Figure 2b); at 45.5 GPa and 750 K, the diffraction spots belonging to the ε -type structure disappeared and a few new spots appeared, implying the occurrence of a phase transition (Figure 2c). This new phase is stable up to 55.0 GPa and 750 K (Figure 2d), the highest *P*-*T* conditions achieved in run-3. When this phase was decompressed to 25.1 GPa at 300 K, it transformed back into the initial ε -type phase (Figure 2f), demonstrating that the detected phase transformation was reversible. The observed new spots could be indexed by a hexagonal structure with lattice parameters a = 6.584(1) Å, c = 4.278(2) Å and V = 160.62(9) Å³ at 44.1 GPa and 300 K (Table S3). These values are very similar to those of hexagonal Fe₇C₃ and Fe₇N₃ (space group: *P*6₃*mc*, denoted as *h*-type) at equivalent *P* (Chen et al., 2012; Minobe et al., 2015). Therefore, an *h*-type structure was selected to

E-type F F F F F F F F F F F F F	ε-type τ τ τ τ τ τ τ τ τ τ τ τ τ	h-type	
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Figure 2. Representative measured 2D XRD patterns of $Fe_7(N_{0.75}C_{0.25})_3$ at high *P-T* conditions in run-3. The ε -type and *h*-type represent the $P3_12$ and Fe_7C_3 -type ($P6_3mc$) structures, respectively.



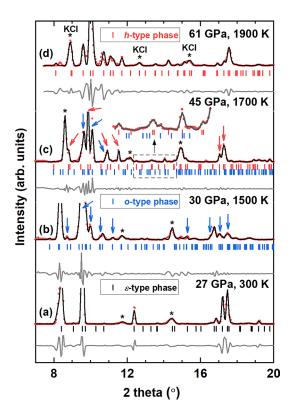


Figure 3. Representative measured XRD patterns of $Fe_7(N_{0.75}C_{0.25})_3$ at high *P-T* conditions in run-4. Le Bail refinements (black solid curves) of observed XRD data (red dots) were carried out after background subtraction. The difference curves (gray continuous line) between observed and calculated spectra are shown. The stars represent Bragg peaks of KCl. The black, blue, and red ticks represent Brag peaks of ε -type (*P*3₁2), *o*-type (Fe₃C-type, *Pnma*), and *h*-type (Fe₇C₃-type, *P*6₃*mc*) phases, respectively.

solve the integrated 1-D XRD pattern, and the refinement results suggest that the new high-*P* phase can indeed be assigned to the *h*-type structure (Figure S5).

In run-4, ε -type Fe₇(N_{0.75}C_{0.25})₃ was compressed to ~27 GPa at 300 K and then laser-heated, and we observed the Bragg peak intensities of the ε -type phase gradually decreased whereas a few new peaks emerged (Figure 3b). The new phase can be assigned to a Fe₃C-type structure (space group: Pnma, denoted as o-type). The refined lattice parameters of the quenched o-type phase were determined to be a = 4.892(2) Å, b = 6.525(2) Å, c = 4.366(1) Å and V = 139.38(4) Å³ at 28 GPa and 300 K (Figure S6b), values that are slightly higher than those of pure Fe₃C at the similar P (Prescher et al., 2012). The XRD pattern after quenching was similar to that during heating, and no Bragg peaks of the ɛ-type structure were observed (Figure S7). When laser-annealed at \sim 40 GPa, the *o*-type and *h*-type phases coexisted, and the *h*-type phase was dominant under these conditions (Figure 3c and Figure S6c). After being laser-annealed at higher P, ~60 GPa, the initial ε -type Fe₇(N_{0.75}C_{0.25})₃ transformed directly into the *h*-type phase without undergoing an intermediate *o*-type phase (Figure 3d and Figure S6d).

DFT-based first-principles calculations were performed to investigate the phase transitions that were observed in run-3 and run-4 (Table S4). We first looked into the magnetic properties of the ε -type, *o*-type and *h*-type structures for the Fe-N, Fe-C and Fe-N-C systems. For the ε -type structure, the magnetic moments of Fe are slightly affected by the substitution of C for N, whereas they are anti-correlated with the Fe/(N+C) ratio (Figure S8a), in agreement with previous calculations concerning the Fe-N system (Popov et al., 2015). The magnetic moments of Fe of all calculated ε -type phases decrease smoothly with *P*, indicating there occurs no spin transition up to 50 GPa. For the *o*-type structure, Fe₃N, Fe₃C and Fe₃(N_{0.75}C_{0.25}) have similar average magnetic moments, and the substitution of C for N tends to slightly increase the difference between the magnetic moments of FeI and FeII (Figure S8b–S8c). In particular, the

magnetic moments of Fe in Fe₃C collapse to zero above ~55 GPa, indicating a FM-NM transition of Fe₃C. This is consistent with previous X-ray emission spectroscopy (XES) results (Chen et al., 2018). By contrast, such transition is not observed in the N-rich *o*-type phases. For the *h*-type structure, Fe₇N₃ and Fe₇C₃ have distinct sub-lattice magnetic properties, though they adopt the same crystal structure and have comparable average magnetic moments (Figure S8d–S8e). Specifically, FeI has the largest magnetic moments in Fe₇N₃, whereas FeII does in Fe₇C₃. The average and sub-lattice magnetic moments of Fe decrease gradually with *P* for Fe₇N₃ and Fe₇C₃, and no spin transition is predicted to occur in the calculated *P* range. It should be noted that there is a discontinuity in the evolution of magnetic moments of Fe in Fe₇C₃ at 40~45 GPa. This phenomenon has not been observed by previous experiments such as Mössbauer spectroscopy (MS) or XES (Chen et al., 2012, 2014). However, previous simulations have reported similar results (Litasov et al., 2015). The reason of the difference between experimental and theoretical results remains to be investigated.

The *o*-type Fe₃C phase appears to be the most stable over the *P* range of the calculations (0~65 GPa) (Figure S9), in agreement with previous experimental results (Liu, Lin, et al., 2016). The difference in unit-cell volumes between experimental and calculated results for pure Fe₃C is less than 2%. The ε -type Fe₃N is predicted to undergo phase transitions to the *o*-type phase at 38 GPa and 0 K, then to an assemblage of *h*-type Fe₇N₃ and Fe at 55 GPa and 0 K (Figure 4a). We only calculated the ε -type and *h*-type phases in a Fe₇N₃ stoichiometry, since simulations of *o*-type Fe₇N₃ were unfortunately beyond our computational capability. The ε -to-*h* transition of Fe₇N₃ is predicted to occur at ~57 GPa (Figure 4b). We further considered the *P*-induced phase transition in the Fe-N-C ternary system with a fixed composition of Fe₃(N_{0.75}C_{0.25}). At 0 K, the ε -type Fe₃(N_{0.75}C_{0.25}) transforms into the *o*-type phase at 26 GPa, and further into an assemblage of *h*-type



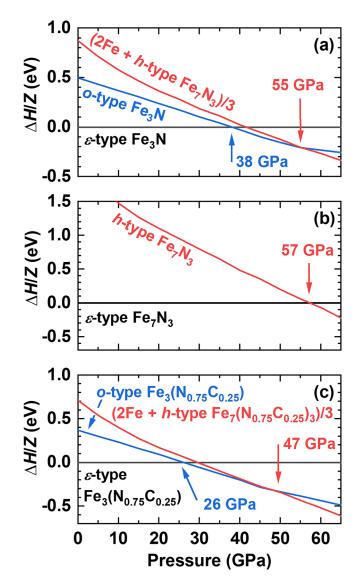


Figure 4. The calculated pressure-enthalpy relations of (a) Fe₃N, (b) Fe₇N₃ and (c) Fe₃(N_{0.75}C_{0.25}) with different structures. The ε -type, *o*-type and *h*-type represent the *P*3₁2, Fe₃C-type (*Pnma*) and Fe₇C₃-type (*P*6₃*mc*) structures, respectively.

 $Fe_7(N_{0.75}C_{0.25})_3$ and Fe at 47 GPa (Figure 4c). The calculated lattice parameters of *o*-type phases are as follows: a = 4.729 Å, b = 6.715 Å, c = 4.323 Å and $V = 137.3 \text{ Å}^3$ for Fe₃N at 29 GPa, and a = 4.760 Å, b = 6.677 Å, c = 4.322 Å and V = 137.4 Å³ for Fe₃(N_{0.75}C_{0.25}) at 28 GPa, respectively. The calculated unit-cell V of o-type Fe₃N is $1.8\% \sim 2.0\%$ larger than that of isostructural Fe₃C (Figure S10a), whereas the unit-cell V of o-type $Fe_3(N_{0.75}C_{0.25})$ lies in between those of Fe_3C and Fe_3N . These results imply that the substitution of N for C can enlarge the unit-cell V of o-type Fe₃C. Furthermore, this substitution tends to reduce the a axis, increase the baxis and slightly affect the c axis (Figure S11b). It will correspondingly reduce the a/c ratio and increase the b/c ratio (Figure S12b). The unit-cell V of the o-type phase detected at 28 GPa and 300 K is $\sim 1.4\%$ larger than that of pure Fe₃C as measured by single-crystal XRD experiments (Figure S10a) (Prescher et al., 2012). The a/c ratio of the o-type phase in our experiments deviates from those of pure Fe₃C. Therefore, it is reasonable to propose that the o-type phase detected in run-4 is an N-enriched phase.

3.2. Thermoelastic Properties of Iron Carbonitrides

The lattice parameters of the two ε -type samples in run-1 and run-2 were determined only to a *P* up to ~33 GPa: Above this *P*, many of the Debye-Scherrer rings or spots associated with the Fe carbonitrides overlap with those associated with Re, the PTM, and Au (Table S5 and S6). The unit-cell volumes of the ε -type phases decrease monotonically with *P*, and no discontinuities in their *P*-*V* curves are observed (Figure 5a). The room-*T P*-*V* data were fitted to the third-order Birch-Murnaghan EoS (Birch, 1947; Murnaghan, 1944):

$$P(V) = \frac{3}{2}K_0 \left[\left(\frac{V_0}{V}\right)^{\frac{7}{3}} - \left(\frac{V_0}{V}\right)^{\frac{5}{3}} \right] \times \left\{ 1 + \frac{3}{4} \left(K'_0 - 4\right) \left[\left(\frac{V_0}{V}\right)^{\frac{2}{3}} - 1 \right] \right\}$$

The normalized stress (F_E) versus Eulerian strain (f_E) plot suggests that it is reasonable to fit the *P*-*V* data using one single equation of state over their respective *P* range (Figure S13). In addition, the slope of the f_E - F_E curve for each ε -type sample is close to zero. The corresponding K_0 ' derived from the f_E - F_E relation is 3.94 ± 0.25 for Fe₄(N_{0.6}C_{0.4}) and 4.15 ± 0.12 for Fe₇(N_{0.75}C_{0.25})₃. These results indicate that it is reasonable to fix K_0 ' as 4, and the *V* compressibility of ε -type Fe carbonitrides could be adequately described by the second-order Birch-Murnaghan EoS. The fitted parameters yield, for ε -type Fe₄(N_{0.6}C_{0.4}), $V_0 = 81.11(7)$ Å³, $K_0 = 133(2)$ GPa,

and $K_0' = 4$ (fixed). For ε -type Fe₇(N_{0.75}C_{0.25})₃, the values are $V_0 = 85.00(1)$ Å³, $K_0 = 177(1)$ GPa and $K_0' = 4$ (fixed) (Table S7). Fe₇(N_{0.75}C_{0.25})₃ has a similar compressibility to pure Fe nitride and is less compressible than Fe₄(N_{0.6}C_{0.4}) (Figure 5b and Table S7) (Adler & Williams, 2005; Litasov et al., 2014, 2017, 2017; Lv et al., 2020; Yin et al., 2014).

The relative compression of crystal axes and the c/a ratio of the ε -type Fe carbonitrides at room T are shown in Figures 5c and 5d. Fe₄(N_{0.6}C_{0.4}) is nearly isotropic in the axial direction and its c/a ratio is almost constant up to at least 30 GPa. The compressibility along the *a* axis of Fe₇(N_{0.75}C_{0.25})₃ is comparable to that along the *c* axis below 20 GPa, but becomes higher than the *c* axis above 20 GPa. This effect is also manifested in a change in the c/a ratio. The c/a ratio of Fe₇(N_{0.75}C_{0.25})₃ is systematically smaller than that of Fe₄(N_{0.6}C_{0.4}). Theoretical calculations show that the incorporation of N or C has a negligible effect on the *c* axis but increases the *a* axis (Figure S11a). And the slope of the c/a ratio versus *P* of Fe₇(N,C)₃ is larger than that of Fe₃(N,C) (Figure S12a). Recent neutron diffraction experiments have found the magnetic moments



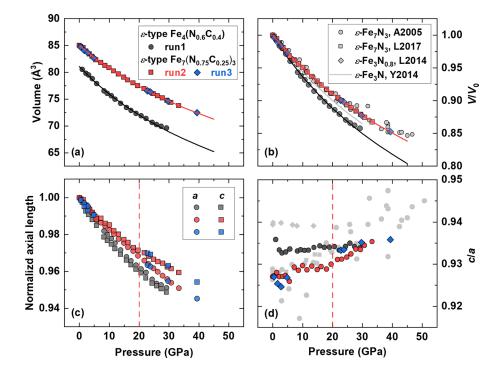


Figure 5. Evolution of (a) unit-cell *V*, (b) normalized *V*, (c) normalized axial length and (d) c/a ratio of ε -type $Fe_4(N_{0.6}C_{0.4})$ and $Fe_7(N_{0.75}C_{0.25})_3$ as a function of *P* at room *T*. The black and red solid curves are calculated using the Birch-Murnaghan equation of state parameters of ε -type $Fe_4(N_{0.6}C_{0.4})$ and $Fe_7(N_{0.75}C_{0.25})_3$, respectively. The red dashed line indicates the abnormal change in the relative axial compression and c/a ratio of ε -type $Fe_7(N_{0.75}C_{0.25})_3$. The uncertainties are within the symbols. Previous experimental data of ε -type Fe nitrides are plotted for comparison (Adler & Williams, 2005; Litasov et al., 2014, 2017, 2017; Yin et al., 2014).

of Fe oriented perpendicular to the *c* axis of the ε -type structure (Lei et al., 2018). The addition of more N or C into the ε -type structure enlarges the *a* axis and reduces the interactions of magnetic moments between Fe atoms, which results in a reduction of magnetic moments (Figure S8). This will cause the *a* axis to become more compressible. The change of the *c/a* ratio of ε -type Fe₇(N,C)₃ is not observed in the present simulation. This change cannot result from the spin transition of Fe based on our calculations. It

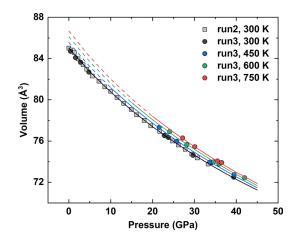


Figure 6. The *P*-*V*-*T* data of ε -type Fe₇(N_{0.75}C_{0.25})₃ fitted by the high-*T* Birch-Murnaghan equation of state. The solid and dashed curves are calculated using the Birch-Murnaghan equation of state parameters of ε -type Fe₇(N_{0.75}C_{0.25})₃. The *P* and *V* uncertainties are within the symbols.

might be related to a change in the magnetic ordering of $Fe_7(N_{0.75}C_{0.25})_3$. However, since our first-principles calculations are performed at 0 K, the paramagnetic state cannot be simulated. Further MS experiments could provide direct evidence on the evolution of magnetic ordering of ε -type Fe carbonitride samples upon compression. Note that even if there is an anomaly in the relative compressibility of the axes, it exerts a negligible effect on the compressive properties of unit-cell *V* of ε -type Fe₇(N_{0.75}C_{0.25})₃ (Figure S13).

The thermal EoS of our sample was derived from its room-*T* EoS and *P*-*V* data obtained at 450 K, 600, and 750 K (Figure 6). Assuming that the *T* dependence of isothermal bulk modulus at high temperature $\partial K/\partial T$ and the first *P* derivative of isothermal bulk modulus K_{T0} ' are constant over the investigated *P*-*T* range, K_{T0} and K_{T0} ' can be expressed as following:

$$K_{T0} = K_0 + \partial K / \partial T \times (T - 300)$$

 $K_{T0}' = K_0'$



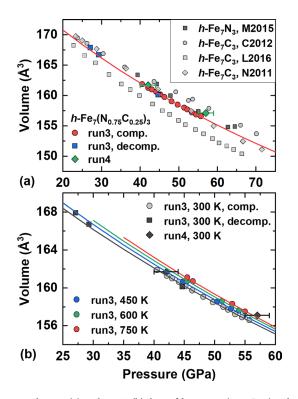


Figure 7. The *P*-*V*(a) and *P*-*V*-*T*(b) data of *h*-type $Fe_7(N_{0.75}C_{0.25})_3$. The solid curves are calculated using the Birch-Murnaghan equation of state parameters of *h*-type $Fe_7(N_{0.75}C_{0.25})_3$. The error bars represent *P* uncertainties in run-4. The *P* and *V* uncertainties in run-3 are within the symbols. Previous experimental data of *h*-type Fe_7C_3 and Fe_7N_3 are plotted for comparison (Chen et al., 2012; Liu, Li, & Ikuta, 2016; Minobe et al., 2015; Nakajima et al., 2011).

The *T* dependence of *V* at ambient *P*, V_{T0} , can be expressed as a function of thermal expansion coefficient α_T :

$$V_{T0} = V_0 exp\left(\int\limits_{300}^{T} \alpha_T dT\right)$$

and due to limited *V* data at simultaneous high *P*-*T* conditions, the thermal expansion coefficient α_T is assumed to be constant by the approximation that $\alpha_T = \alpha_0$ (Berman, 1988). The fitted thermoelastic parameters for ε -type Fe₇(N_{0.75}C_{0.25})₃ are consequently $\partial K/\partial T = -0.02(1)$ GPa/K and $\alpha_0 = 3.3(8) \times 10^{-5}$ K⁻¹ (Table S7).

The refined lattice parameters and unit-cell volumes of h-type $Fe_7(N_{0.75}C_{0.25})_3$ at various P-T conditions are listed in Table S8. The unit-cell volumes of h-type Fe7(N0.75C0.25)3 obtained in run-4 are generally consistent with the results of run-3 (Figure 7a). However, because of the relatively large uncertainty in determining the P in run-4 using KCl, the room-T EoS of h-type $Fe_7(N_{0.75}C_{0.25})_3$ is calculated only using the P-V data during compression in run-3. The fitted parameters for *h*-type Fe₇(N_{0.75}C_{0.25})₃ are $V_0 = 182.4(7)$ Å³, $K_0 = 268(9)$ GPa and $K_0' = 4$ (fixed). The compressibility of *h*-type $Fe_7(N_{0.75}C_{0.25})_3$ is generally comparable to that of the isostructural endmembers Fe₇N₃ and Fe₇C₃ (Figure 7a and Table S9) (Chen et al., 2012; Kusakabe et al., 2019; Liu, Li, & Ikuta, 2016; Minobe et al., 2015; Nakajima et al., 2011). The lattice parameters of $Fe_7(N_{0.75}C_{0.25})_3$ decrease smoothly with increasing P (Figures S11c and S14). The a/c ratio of h-type Fe₇(N_{0.75}C_{0.25})₃ increases slightly with P as compared with Fe₇C₃, implying a weak anisotropy in the axial compressibility (Figure S12c). Although there are some differences between experimental and theoretical data in terms of the absolute values, the calculated results generally follow the trend of the experimental ones. First-principles simulations have shown that N-rich Fe₇(N,C)₃ has distinct sub-lattice magnetic properties from Fe₇C₃. The Fe-Fe distance in the *ab* plane is shorter than that along the *c* axis in the *h*-type structure.

FeII has the largest magnetic moments in Fe₇C₃ but it is the smallest in Fe₇N₃. Therefore, there exists stronger magnetic and electronic interactions between Fe atoms in the *ab* plane in Fe₇C₃ in comparison with Fe₇N₃. This contributes to different axial compression behaviors of *h*-type Fe₇C₃ and N-rich Fe₇(N,C)₃. We also calculated the thermal EoS of *h*-type Fe₇(N_{0.75}C_{0.25})₃ with the following fitted results: $\partial K/\partial T = -0.046$ (34) GPa/K and $\alpha_0 = 3.2 (1.5) \times 10^{-5} \text{ K}^{-1}$ (Figure 7b).

4. Discussion

Previous experiments have reported that ε -type Fe₇N₃ transforms directly into the *h*-type structure above 40 GPa at high *T*, and the *o*-type phase has not been observed between 20 and 40 GPa (Minobe et al., 2015). Our static calculations have shown that *o*-type Fe₃N becomes energetically favorable in comparison with ε -type Fe₃N above 38 GPa. However, the ε -to-*o* transition is only evaluated in the Fe₃X stoichiometry due to our limited computational capability. The *o*-type Fe₃C is found to have a wide solid solution range depending on the synthetic conditions. The *o*-type Fe₃C with C-deficiency up to 2.6 at% (Fe₃C_{0.87}) is synthesized at Fe-saturated conditions, and that containing 28 at% C (Fe₃C_{1.17}) is detected at C-saturated conditions (Walker et al., 2013, 2015; Zhu, Li, Walker, et al., 2019). Adding much more C to stoichiometric Fe₃C tends to destabilize the *o*-type phase and finally leads to forming *h*-type Fe₇C₃. If this can be applied in the Fe-N system, the present calculations may overestimate the stability of the *o*-type Fe nitride. The incorporation of C in the Fe-N system can lower the two transition pressures and slightly expand the stability field of the *o*-type phase, which may facilitate the observation of this phase in Fe carbonitrides. However, given the structural characteristics of the *o*-type phase, the ε -to-*o* transformation of Fe₇(N_{0.75}C_{0.25})₃ will be accompanied by a



partial dissociation, causing its composition to approach the Fe₃X stoichiometry. Recent studies have found that deviation from the stoichiometric composition of Fe₃C in either C-deficiency or C-excess direction reduces the unit-cell *V* (Walker et al., 2013, 2015; Zhu, Li, Walker, et al., 2019). This finding indicates that the *o*-type phase observed in the present study is a N-enriched non-stoichiometric solid solution with the chemical formula of Fe₃(N,C)_{1+x} (x < 0.28). In addition, there is no evidence of back-transformation to the ε -type structure from the *o*-type phase after quenching at 27 GPa. Both experimental and theoretical results demonstrated that the ε -to-*o* transition introduces a considerable *V* reduction (Figure S10a). The *o*-type structure should be a high-*P* phase of the ε -type structure instead of a high-*T* one. The *P* of ~27 GPa should be the upper limit of the ε -to-*o* transition boundary, and ε -type Fe₇(N_{0.75}C_{0.25})₃ is metastable in the *P* of 27~40 GPa and in the *T* of 450~750 K (Figure 1).

The *h*-type phase is experimentally detected above ~40 GPa for either Fe nitride or Fe carbonitride, implying that the incorporation of C in the Fe-N system affects the stability of the *h*-type phase slightly. In the Fe-C system, Fe₃C melts incongruently to form Fe₇C₃ plus liquid above 10 GPa (Fei & Brosh, 2014). However, Fe₃C is stable over a broad *P*-*T* field in the solid state, and decomposes into Fe₇C₃ plus Fe above 145 GPa prior to melting (Liu, Lin, et al., 2016). Subsequent experimental studies suggest that Fe₃C and Fe coexist at subsolidus conditions to at least 203 GPa (Mashino et al, 2019). Although the exact stability field of Fe₃C at multi-megabar pressures remains controversial, previous results support the conclusion that Fe₃C cannot decompose into Fe₇C₃ plus Fe below the solidus at pressures below one megabar, a conclusion also supported by first-principles calculations (Mookherjee et al., 2011). Therefore, the incorporation of N in the Fe-C system can stabilize the *h*-type phase at high *P*-*T* conditions. It is expected that C-rich Fe₃(C,N) is able to dissociate into Fe₇(C,N)₃ plus Fe below the solidus at pressures much lower than that for pure Fe₃C.

Three kinds of N-enriched solid solutions, that is, ε -type, o-type and h-type phases, have been experimentally observed. The *V* changes caused by the substitution of C for N are small, generally less than 2% in these solid solutions (Figure 7a and Figure S10a), indicating that the molar volumes of N and C in these structures are very similar. Sokol et al. (2017) found that the solubility of N in o-type Fe₃C is limited to 0.5 wt% at 7.8 GPa, and that the incorporation of 0.5 wt% N leads to a ~10% *V* expansion. This *V* change is inconsistent with the present calculations. One possible reason might be that o-type Fe₃(C,N)_{1+x} is unstable on the C-rich side at ambient conditions, and the unit-cell *V* expands drastically during *P* quenching. More studies are required to reconcile the discrepancy between experimental and theoretical results. Nevertheless, a large *V* expansion may be responsible for impeding N solubility in o-type Fe₃C (Sokol et al., 2017). These results may further imply that Fe₃N and Fe₃C cannot form a complete solid solution in the ε -type or o-type phase at low *P*, since they have distinct crystal structures (Figure S1). The N-enriched ε -type Fe₃(N,C)_{1+x} transforms into the o-type phase at 26~30 GPa, suggesting a possible complete solid solution at 25~40 GPa across the o-type Fe₃C-Fe₃N join. Similarly, *h*-type Fe₇C₃ and Fe₇N₃ are likely to form a complete solid solution in the stability field of the *h*-type phase (i.e., >40 GPa), as observed in diamond inclusions from the lower mantle as well as in our experiments (Kaminsky & Wirth, 2017).

At room T, no structural transition occurs in Fe nitrides or carbonitrides (including cubic γ' -Fe₄N, orthorhombic ζ -Fe₂N and ε -type Fe₃(N,C)_x) at pressures up to at least ~70 GPa (Adler & Williams, 2005; Lv et al., 2020; Zhuang et al., 2021). By contrast, previous and present theoretical calculations show that these compounds become unstable at pressures much lower than 70 GPa at 0 K (Niewa, Rau, Wosylus, Meier, Wessel, et al., 2009; Schwarz et al., 2009). Indeed, at high temperatures, these compounds transform into high-P polymorphs at significantly lower pressures. For instance, γ' -Fe₄N and ζ -Fe₂N transform into ε -type Fe₄N and Fe₅N at ~10 GPa at high temperatures, respectively (Guo et al., 2013; Schwarz et al., 2009). The ε -type Fe nitrides undergo a reversible structural transition to the *h*-type phase above ~40 GPa when laser-annealed (Minobe et al., 2015). The discrepancies between theoretical prediction and experimental observations suggest a kinetic barrier in the structural transitions for Fe nitrides and carbonitrides. This hinders the transformation to high-P phases, and consequently renders the low-P phases metastable to much higher pressures. High T considerably contributes to the rearrangement of atoms to form high-P polymorphs for Fe nitrides and carbonitrides. Based upon our experimental results, the heating T of 750 K is sufficient to overcome the kinetic barrier of the ε -to-h transition for Fe₇(N_{0.75}C_{0.25})₃. In the meanwhile, the kinetic barrier of the *z*-to-*o* transition is still unclear. So far, we can only obtain that the kinetic barrier of this transition cannot be overcome by heating to 750 K but can be overcome by heating to \sim 1,200 K.



The ε -to-h structural transition observed in run-3 introduces a ~4% V collapse in Fe₂(N_{0.75}C_{0.25})₃ at ~40 GPa (Figure S10b), which is close to the value (\sim 6%) for pure Fe₇N₃ (Minobe et al., 2015). The V change for Fe₃N or Fe₃(N_{0.75}C_{0.25}) is theoretically determined to be ~6% across the ε -to-o structural transition (Figure S10a). The ε -type structure is derived from a distorted hexagonal close-packed (*hcp*) structure of Fe. N (or C) atoms occupy the octahedral sites in an ordered or disordered way with many unoccupied vacancies (Niewa, Rau, Wosylus, Meier, Hanfland, et al., 2009). In o-type and h-type structures, each N (or C) atom is coordinated to six Fe atoms, forming a ditrigonal prism (Figure S1). The difference in the arrangement of atoms is responsible for a considerable V reduction through the ε -to-o or ε -to-h structural transition. The experimentally determined K_0 of ε -type Fe nitrides and carbonitrides varies from ~130 to ~180 GPa (Table S7) (Adler & Williams, 2005; Litasov et al., 2014, 2017; Lv et al., 2020; Yin et al., 2014). Our experimental and theoretical results suggest that the substitution of C for N has only a slight effect on the compressibility of the ε -type phase (Figure 5 and Table S4). In contrast, K_0 of the ε -type phase strongly depends on the total concentrations of N and C. With more N and C in the ε -type phase, fewer unoccupied octahedral sites will be present. The NFe₆ (or CFe₆) octahedron is less compressible in comparison to the octahedral vacancy. Therefore, K_0 of the ε -type phase is negatively correlated to the ratio of Fe/(N+C). In terms of the h-type phase, if K_0 is fixed to 3.2 as reported by Chen et al. (2012), the corresponding K_0 are 306 (9) GPa, 307 and 316 GPa for $Fe_7(N_{0.75}C_{0.25})_3$, Fe_7C_3 and Fe_7N_3 , respectively (Table S9) (Chen et al., 2012; Kusakabe et al., 2019). With a fixed Fe/(N+C) ratio of ~2.33, the K_0 of the *h*-type phase is nearly independent of the N/C ratio. The thermoelastic parameters of ε -type Fe₇(N_{0.75}C_{0.25})₃ obtained in the present study (see Table S7) are similar to those of ε -type Fe₃N_{0.8} and Fe₇N₃ (Litasov et al., 2014, 2017). Moreover, both *h*-type Fe₇(N_{0.75}C_{0.25})₃ and Fe₇C₃ have comparable thermoelastic properties (Lai et al., 2018). To summarize, we suggest that for the Fe-N-C system, the stability and phase transitions of Fe carbonitride at high P-T conditions are mainly controlled by the N/C ratio, with the Fe/(N+C) ratio significantly affecting its isothermal compressibility, whereas the compositional effect plays a negligible role in influencing its thermoelastic parameters (i.e., $\partial K/\partial T$ and α_T).

5. Implications

A recent study has estimated that more than 80% of N in the slab does not return to the surface in the cold slab geotherm (Busigny et al., 2019). Even in the warm slab, about half of N in the slab can be transported to the deep Earth. Potassium silicates, such as cymrite, wadeite and hollandite, are major carriers transporting N in the lithosphere to the lower upper mantle and transition zone via subduction (Sokol et al., 2020; Watenphul et al., 2009). Stishovite, as one of major phases in subducted slabs, is shown to be capable of incorporating a few hundreds of ppm N at 28 GPa and 1,700~2,000 K, indicating that a large amount of N could have been continuously transported into the lower mantle to at least ~800 km through subduction (Fukuyama et al., 2020). Recent petrologic observations and mineral physics experiments suggest that native Fe and Fe-enriched alloys may exist not only at the core-mantle boundary (CMB), but also in the bulk deep mantle (Armstrong et al., 2019; Frost et al., 2004; Smith et al., 2016). The Fe-enriched domains in the deep mantle may be alloyed with N throughout geological times, provided that N is continuously supplied by the slabs. The interactions between subducted slabs and Fe-enriched domains can lead to the formation of Fe-N or Fe-N-C compounds during the slab-mantle interaction (Figure 8). The structures of N-bearing Fe-enriched phases depend on the formation depth and N/C ratios. In general, ε -type and o-type phases are dominant above \sim 1,000 km, and they will both transform into the *h*-type phase below \sim 1,000 km. Nitrogen is likely to be residing in the deep mantle in the form of Fe nitrides and carbonitrides with diverse polymorphs and compositions. This may keep N hidden from observations at the surface and partially solve the "missing nitrogen" problem. In addition, the Fe-enriched phase can partially or fully melt with the presence of sulfur (S) or nickel (Ni) at the level of several weight percent (Gilfoy & Li, 2020). The solubility of N in the Fe-enriched melt remains nearly unaffected in the metallic melt containing up to ~17 wt% S (Grewal, Dasgupta, Holmes, et al., 2019). Therefore, the Fe-enriched melt can still efficiently extract N from the slabs even when other volatiles are present. Furthermore, there is a large difference in the N bonding between metal alloys (ε -type, o-type and h-type phases) and subducted slabs. The extraction of N from the slabs to the Fe-enriched domains in the bulk mantle is anticipated to be accompanied by considerable N-isotopic fractionations, which can be verified by future studies.



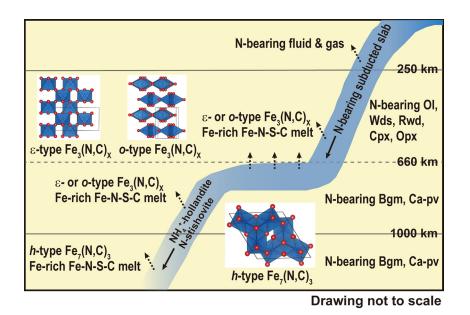


Figure 8. Schematic model for possible N hosts in the deep mantle. The N stored in the mantle can be either from the residue of the primordial mantle or from the injection of N via subducted slabs. Nitrogen can be stored in major silicates of the normal mantle. Below 250 km, where the mantle is saturated with metallic Fe, ε -type and *o*-type Fe nitrides and carbonitrides become important hosts for N. The *h*-type Fe nitrides and carbonitrides are energetically favorable in comparison to ε -type and *o*-type counterparts below 1,000 km, and they should be the most significant metallic host for N down to at least 1,500 km.

Superdeep diamonds carry unique information about the mineralogy and geochemistry of the inaccessible deep Earth (Kaminsky, 2012). Most superdeep diamonds are typically N-free or only contain N concentrations at the level of several ppm (Smith et al., 2016). However, recent studies have detected diverse Fe nitrides, carbonitrides, nitrocarbides and carbides as inclusions in lower-mantle diamonds (Kaminsky & Wirth, 2017). The compositions of Fe-N-C inclusions in superdeep diamonds are complex, with the Fe/ (N+C) ratio ranging from 1.67 to 3.83 and the N/(N+C) ratio ranging from 0.2 to 0.8. Of particular interest is the fact that Fe_7C_3 and ε -type $Fe_9(N_{0.8}C_{0.2})_4$ coexist as inclusions in a lower-mantle diamond (Kaminsky & Wirth, 2017). Based on the compositional and structural analysis of two inclusions and the phase relation of the Fe-C system at high pressures, Kaminsky and Wirth (2017) suggested that the observed Fe alloys probably originated from the CMB after migrating from the Earth's liquid outer core. Although Fe and Fe_7C_3 become eutectic at very high pressures, Fe_7C_3 can be formed by the incongruent melting of Fe_3C above \sim 10 GPa (Fei & Brosh, 2014; Liu, Lin, et al., 2016). In addition, Fe₇C₃ can occur as an intermediate by-product of the reaction between metallic Fe and subducted carbonates (Zhu, Li, Liu, et al., 2019). This assemblage likely represents reduced domains in the lower mantle, which can be locally preserved throughout the mantle and contribute to the primordial redox heterogeneity of the mantle. However, the occurrence of this assemblage as inclusions may not necessarily indicate a CMB origin (Zedgenizov & Litasov, 2017). Another possibility is that this assemblage is formed through a two-stage crystallization process involving a Fe-N-C melt in the so-called metallic pockets in the deep mantle (Smith et al., 2016). Fe₇C₃ has the highest melting temperature among Fe carbides and nitrides, and it will crystallize first from the melt (Kusakabe et al., 2019; Liu, Lin, et al., 2016). However, the TEM image showing Fe_7C_3 transects $Fe_9(N_{0.8}C_{0.2})_4$ indicates that $Fe_9(N_{0.8}C_{0.2})_4$ is formed no later than Fe_7C_3 . In addition, this model cannot explain why Fe_7C_3 coexisting with N-rich $Fe_9(N_{0.8}C_{0.2})_4$ is totally N-free, unless N only partitions into metallic melt when in equilibrium with Fe₇C₃. This explanation is contradicted by the observation of a Fe₇($C_{0.73}N_{0.27}$)₃ inclusion in a diamond from the lower mantle (Kaminsky & Wirth, 2017).

Based on our results, we propose that the origin of this assemblage is an *h*-type N-enriched Fe₇(N,C)₃ precursor phase. When *h*-type Fe₇(N,C)₃ is carried into shallow regions, it will transform into the low-*P* ε -type phase as observed in the present study. The incorporation of C in the ε -type phase has an upper limit that is controlled by *P* and *T* (Sokol et al., 2017). This phase transformation will be accompanied by the exsolution



of Fe₇C₃, which can be described as *h*-type Fe₇(N_{1-m}C_m)₃ $\rightarrow \varepsilon$ -type Fe₇(N_{1-m}C_n)₃ + Fe₇C₃ (m > n); this possibility is supported by the fact that Fe₇C₃ and ε -type Fe₉(N_{0.8}C_{0.2})₄ have similar Fe/(C+N) ratios (Kaminsky & Wirth, 2017). The N-enriched Fe₇(N,C)₃ precursor can either be the residue of the primordial mantle or be formed by the interaction between subducted slabs and Fe-enriched domains in the lower mantle. The isotopic characteristics of inclusions in superdeep diamonds can in principle place additional constraints on the origin of these light elements (Dalou et al., 2019). On the other hand, the *o*-type Fe carbonitride or nitrocarbide has not been found in superdeep diamonds yet. The *o*-type phase observed in the present study can be *T*-quenched, but it remains unclear whether or not it is quenchable to ambient *P*. Further investigations are required to constrain the stability field of the *o*-type phase as a function of its N/C ratios. In addition to silicates and oxides, the structural and compositional features of Fe-N-C inclusions as a function of *P*, *T* and even redox conditions can also shed light on the depth of formation of superdeep diamonds and geodynamic processes in the lower mantle.

Recent studies have proposed that the Earth's inner core can consist predominantly of h-type Fe_7C_3 or Fe_7N_3 because their densities can match the density of the inner core (Chen et al., 2012; Kusakabe et al., 2019). The *h*-type $Fe_7(C,N)_3$ is also suggested to be a candidate constituent of the inner core under the assumption tion that its two endmembers could form a continuous solid solution at high P-T conditions (Kusakabe et al., 2019; Minobe et al., 2015). Our study provides the first experimental confirmation of the existence of an *h*-type Fe_7C_3 - Fe_7N_3 solid solution above 40 GPa. We also find that the substitution of C for N slightly affects the compressibility and thermoelastic properties of Fe-N-C compounds with a fixed Fe/(N+C) ratio. It is expected that the density of an intermediate h-type $Fe_7(C,N)_3$ phase, such as $Fe_7(N_{0.75}C_{0.25})_3$ in the present study, will be very close to that of its two endmembers when extrapolated to much higher P-T conditions. On the other hand, a recent theoretical study has predicted a new monoclinic high-P phase for Fe_7N_3 and Fe_7C_3 above ~250 GPa (Sagatov et al., 2019). Subsequent simulations suggested that Fe_7C_3 might decompose into the mixture of Fe₂C and Fe₃C at conditions of the Earth's inner core (Sagatov et al., 2020). Whether these results could be applied in the Fe-N-C ternary system is still under debate. More experimental and theoretical studies are required to determine the phase stability, density and physical properties of h-type $Fe_7(C,N)_3$ solid solutions at simultaneous high P-T conditions relevant to the Earth's core, which can offer more clues to decipher the Earth's core compositions and the speciation of N in the Earth's core.

Data Availability Statement

All the data necessary to produce the results are available on Zenodo (http://doi.org/10.5281/ zenodo.4556877).

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